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MOTION AND IMPACT OF ICEBERGS

DEVELOPMENT OF A MODEL TO PREDICT ICE MASS MOTIONS IN THE VICINITY OF AN OFFSHORE STRUCTURE

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MOUVEMENTS ET IMPACT DES ICEBERGS

REALISATION D'UN MODELE POUR LA PREVISION DES MOUVEMENTS DES MASSES GLACIAIARES A PROXIMITE D'UNE STRUCTURE OFFSHORE

1. RESUME

La firme Hay and Company Consultants Ltée a été retenue par le Fonds Renouvelable pour l'etude de l'Environnement en vue d'étudier les mouvements des masses glaciaires à proximité d'une structure offshore. L'ensemble de l'étude a été divisé en les points suivants:

- 1. tour des documents existants touchant à la question;
- revue générale des modèles informatiques existants utiles pour l'étude du problème;
- 3. réalisation et essai d'un modèle informatique adapté à l'étude des interactions hydrodynamiques de deux corps soumis à l'action des vagues et des courants;
- 4. revue générale des installations équipées de bassins et pouvant conduire des essais sur le comportement des masses glaciaires à proximité immédiate d'une structure.

Le tour des documents informatiques existants n'a permis de receuillir que très peu d'information directement utilisable en vue d'une étude générale de la question. Un nombre considérable de recherches ont été menées sur les problèmes connexes et composants: dérive des masses glaciaires en eaux libres et mouvements, provoqués par les vagues, d'un corps flottant librement en eaux libres. Les documents relatifs à ces problèmes connexes sont analysés dans le présent rapport.

La revue générale des modèles informatiques existants a prouvé que la plupart des modèles appropriés partent due principe de la diffraction d'ondes. Le modèle qui a été retenu dans la présente étude en vue d'une réalisation plus poussée est celui de M.M. Isaacson, qui était l'un des chercheurs principaux chargés de cette étude.

Il a été réalisé un modèle permettant de prévoir les mouvements des masses glaciaires sous l'effet des vagues et des courants à proximité d'une forte structure fixe située au large. utilisée repose sur la superposition du mouvement oscillatoire provoqué vagues et obtenu par par des la theorie diffraction d'ondes linéaires, d'une part, et des mouvements de dérive calculés au moyen d'une procédure séquentielle, Cette méthode est appliquée à deux corps voisins de forme tri-dimensionnelle arbitraire, dont l'un est fixe (la structure) l'autre libre et flottant, avec six degrés de liberté glaciaire). Les effets d'interaction hydrodynamique produits entre les deux corps sont inclus dans la procédure, ceux d'amortissement visqueux sont pris en compte l'inclusion de coefficients d'amortissement empiriques. Pour le calcul des mouvements de dérive, on considère que la glaciaire est affectée de deux degrés de liberté (déplacements vertical et latéral). Les équations du mouvement comprennent l'addition de masses de fréquence zéro, les forces de (vagues) et de traînée (courants), tous facteurs qui varier en fonction de l'emplacement de la masse glaciaire eu égard à la proximité de la structure.

Une analyse partielle paramétrique et de sensibilité a été conduite en vue de démontrer les possibilités de ce modèle. Dans les passages d'essais, la structure est représentée par un gros cylindre circulaire vertical et qui touche au fond, et trois sortes de masses glaciaires sont considérées qui correspondent toutes à des cylindres circulaires en flottaison. Les essais comportent des variations des facteurs suivants: volume de la masse glaciaire, direction des vagues, emplacement initial de la masse glaciaire, profondeur de l'eau et force du courant, et ils permettent d'étudier l'influence des forces de dérive (vagues).

Les résultats sont présentés de façon à illustrer la trajectoire de la masse glaciaire et à montrer les variations de sa vélocité que les coefficients hydrodynamiques particuliers à chaque Ils révèlent que, pour les profils géométriques considérés, l'addition de masses de fréquence zéro est tout à insensible à la proximité de la masse glaciaire et structure, sauf dans les cas de proximité immédiate, et que la force de dérive (vagues) influe de façon plus sensible sur trajectoire de la dérive des masses glaciaires de faible volume.

La revue générale des installations disposant de bassins a permis l'évaluation de seize laboratoires équipés d'installations où peuvent se conduire des études dans le domaine en question. Les détails touchant aux dimensions respectives de ces installations et à leur capacité de générer vagues et courants ont été classifiés. Les laboratoires sont évalués et de plus des idées sont proposées sur la façon de conduire les essais.

MOTION AND IMPACT OF ICEBERGS

DEVELOPMENT OF A MODEL PREDICTING ICE MASS MOTIONS IN THE VICINITY OF AN OFFSHORE STRUCTURE

1. SUMMARY

Hay and Company Consultants Inc. have been engaged by the Environmental Studies Revolving Funds to investigate ice mass motions in the vicinity of an offshore structure. The scope of the work was divided into the following tasks:

- a review of available literature pertaining to the problem;
- a survey of existing computer models useful in investigating the problem;
- 3. development and testing of a computer model appropriate for the hydrodynamic interactions of two bodies in waves and currents;
- 4. a survey of model tank facilities capable of performing tests of the behaviour of ice masses in close proximity to a structure.

The literature review identified very little information directly applicable to the overall study problem. A considerable number of studies have been carried out on the related and component problems dealing with ice mass drift in open water and with wave-induced motions of a freely floating body in open water. The literature on these related problems is discussed in this report.

The survey of existing computer models determined that the most appropriate models are those based on wave diffraction. The model utilized in this study for further development was that of

Dr. M. Isaacson, a principal investigator for this study.

A numerical model was developed for predicting ice mass motions due to waves and currents in close proximity to a large fixed offshore structure. The method is based on superposing the waveinduced oscillatory motion obtained by linear wave diffraction theory on the drift motions computed through a time-stepping The method pertains to two neighbouring bodies procedure. arbitrary three dimensional shape. One body is fixed (the structure) and the other is freely floating with six degrees of freedom (the ice mass). Hydrodynamic interaction effects between the two bodies are included in the procedure. Viscous damping effects are taken into account by the inclusion of empirical damping coefficients. In the computation of drift motions, the ice mass is considered to have two degrees of freedom (surge and The equations of motion include zero frequency added wave drift forces and current drag forces, all of which may vary with ice mass location owing to the proximity of the structure.

A limited parametric and sensitivity analysis was carried out to demonstrate the model's capabilities. In the test runs, the structure is represented as a large, vertical, circular cylinder extending to the seabed. Three different ice masses are modelled, all corresponding to floating circular cylinders. The tests include variations in ice mass size, wave direction, initial ice mass location, water depth and current magnitude and examine the influence of wave drift forces.

Results are presented showing the trajectory of the ice mass, and the variations of ice mass velocity and selected hydrodynamic coefficients for each test. These results indicate that, for the geometries considered, the zero frequency added mass is quite insensitive to the proximity of the ice mass to the structure, except when it is very close to the structure, and that the wave drift force has a significant influence on the drift trajectory

of smaller ice masses.

The survey of model tank facilities assessed sixteen laboratories with facilities capable of studies in the problem area. The particulars of these facilities relative to their size, wave generating capabilities and current generating capabilities have been tabulated. The laboratories are rated and suggestions made regarding testing methods.

2. CONCLUSIONS AND RECOMMENDATIONS

- 2.1 Prior to this study, very little information was available regarding the drift motions of an ice mass due to waves and currents in close proximity to a large fixed offshore structure.
- 2.2 It was determined that existing computer models possessing a wave diffraction program would be suitable as a starting point for developing a model predicting drift in the near vicinity of an offshore structure. A model developed by Dr. M. Isaacson was utilized in this study.
- 2.3 A computer model has been developed for predicting ice mass motions in the near vicinity of an offshore structure. The model has been demonstrated to perform satisfactorily for a range of input conditions. These have been chosen to reflect variations in water depth; ice mass size; wave direction; current magnitude and direction; initial location and velocity of the ice mass; and degree of discretization of the ice mass and structure.
- 2.4 Small ice masses experience a negative drift force for certain ranges in the distance from the structure. This is consistent with previous results carried out for two body systems. A significant consequence of this variation in wave drift force is that the drift motion reverses.
- 2.5 The zero frequency added masses change significantly only when the gap between the ice mass and structure is very small. Further improvements to the routine used to calculate the Green's function for the zero frequency added mass determination are recommended.
- 2.6 The case of reduced depth, with a gap of 10 m between the ice mass and seabed, showed no significant effect on the ice

mass motions, bearing in mind that the wave length was altered in this study due to reducing the water depth to attain the 10 m gap.

- 2.7 Only a limited parametric and sensitivity analysis was undertaken in conjunction with the model development. A more comprehensive series of numerical tests should be carried out in order that additional parameters are varied and the influence of these examined.
- 2.8 The computer model may readily be extended to random waves although further numerical improvements in the diffraction calculation would be required to make it computationally efficient.
- 2.9 By incorporating the relevant ice properties, the computer model may ultimately be extended to predict ice mass motions after impact and the resulting forces on a structure.
- 2.10 A series of physical model tests should be carried out and used as a basis for verifying and calibrating the computer model developed in this study.
- 2.11 The combined use of physical and computer models would reduce testing costs in a laboratory study and would be particularly useful in determining cause effect relationships in observed model behavior.
- 2.12 Consideration should be given to towing the fixed structure as a means of creating a uniform current distribution. Careful attention would have to be taken in minimizing the distortion of waves.
- 2.13 On technical considerations, SSPD Maritime Research and Consulting in Sweden ranked highest in the tank survey. The David Taylor Naval Ship Research and Development Center in

the USA and Marintek A/S in Norway ranked a very close second while Hydraulic Research Ltd. and British Maritime Technology in England and Maritime Research Institute in Holland ranked a very close third. Any of the above facilities should be actively considered in a laboratory study.

3. INTRODUCTION

Hay and Company Consultants Inc. have been engaged by the Environmental Studies Revolving Funds of the Government of Canada to undertake a study of the motion and impact of ice masses the vicinity of large, fixed offshore structures. In the last several years, exploration drilling has been carried out on large scale off Canada's east coast. Iceberg encroachment in the Labrador Sea, the Flemish Pass and on the Grand Bank poses unique problem for drilling rigs and for future production Presently, collisions are avoided by a variety of facilities. including long range surveillance and moving threatened rigs. As production facilities under consideration bottom founded structures, impacts with icebergs must considered in their design.

The general problem involving impact from floating ice masses on a large fixed offshore structure covers a wide range of possible parameters. Typical ice masses range in diameter from under 10 m to more than 200 m, and in mass up to about 10 tonnes. East Coast offshore activity generally extends to water depths up to about 100 m but may extend to about 300 m. Structures of approximately 100 m diameter have been proposed for the area.

The original Request for Proposal involved a two component study:

- (1) investigation of methods to estimate the drift and instantaneous velocities induced in an iceberg of known size and shape under known environmental conditions.
- (2) investigation of methods to estimate the behaviour of icebergs in close proximity to a large, fixed offshore structure.

As a basis for a terms of reference, these components are rather broad, and could encompass both the drift of an iceberg in the open ocean over long periods of time, and the drift of an iceberg in the near vicinity of an offshore structure. While these two problems are related, each requires a different computer model to properly address the phenomena involved.

At the initial meeting with the ESRF scientific advisor, it was decided to concentrate on the ice motions (including both drift and oscillatory components) in the near vicinity of an offshore structure.

Iceberg/structure impact is a complex phenomemon, and aspects of it have been the subjects of intensive investigation. An impact event may be divided into three separate phases or regions:

Far-Field Phase:

The ice mass is sufficiently distant from the structure to be unaffected by it. Therefore, the ice mass motion is purely in response to the environmental driving forces.

2. Near-Field Phase:

The ice mass and structure are in close proximity, and the far field motions are altered in response to the structure's presence.

3. Contact Phase:

The ice mass and the structure are in physical contact. The impact is governed by the material and mechanical properties of both, the initial impact conditions, the hydrodynamic interactions and the environmental forces.

Considerable effort has been expended addressing aspects of Phases 1 (quantifying environmental driving forces, iceberg drift, iceberg occurrence and mass distributions), and 3 (quantifying ice material and mechanical properties, iceberg geometries, structure-foundation response). Little has been done

directly in addressing Phase 2.

Two aspects of the near field phase have immediate application in assessing impact loads. First, the effective kinetic energy that ice mass brings to the physical impact may be appreciably altered from its far-field value. The change in energy result from changes induced by the presence of the structure the velocity and added mass. Secondly, the presence structure may alter the eccentricity of the impact, iceberg trajectory changes in response to the near-field flow. eccentricity modifies not only the magnitude but also the torsional load and the overall frequency of impacts occurring.

The near-field phase itself may be divided into two regions:

2a. Inertial Region:

The ice mass and structure are in close proximity, but are sufficiently separated so that the ice mass is only influenced by changes to the far-field flow induced by the structure. Such changes include diffraction and reflection of the wave field and distortion of the current.

2b. Viscous (or Very Near-Field) Region:

As the distance between the ice mass and structure decreases, the fluid viscosity becomes an increasingly important, and perhaps dominant, factor. The very near-field is that range of separations in which fluid friction cannot be ignored.

The main thrust of this study is directed to Region 2a, the inertially-dominated, near-field phase. To a lesser degree, far-field topics are addressed in that the response of ice masses in open water to waves and currents is of interest. Detailed investigation of the ice mass behaviour in Region 2b, and the

inclusion of hydrodynamic influences during Phase 3 are beyond the scope of the present study.

The specific tasks performed under this study are:

- (a) Conduct a literature and data review and include not only information directly applicable to the problem under investigation but also general drift of floating bodies in the vicinity of a structure (Section 4).
- (b) Identify available analytical models and evaluate the applicability of these models to the investigation (Section 5).
- (c) Starting with one of the models identified in (b) above, develop a model capable of predicting drift of an iceberg in the vicinity of an offshore structure (Sections 6 and 7).
- (d) Undertake a survey of available model tank facilities suitable for studies of iceberg motions in the vicinity of an offshore structure (Section 8).

4. LITERATURE AND DATA REVIEW

4.1 General

Estimation of ice mass motions, particularly in the vicinity of a fixed structure, has received only limited attention. of related, more thoroughly investigated, topics are applicable to the problem. These topics include wave loads structures, response of floating bodies to waves, interference effects between floating bodies and nonlinearities excitation and structure response. Due to obvious practical applications, literature on floating bodies is particularly voluminous. Therefore, the literature and data review, logically began with the limited information on iceberg response to waves and with the more substantial work on iceberg drift open water, expanded to include these diverse topics. section summarizes the pertinent work in each area. Under each topic, various avenues have been developed, and most do directly relate to the method developed in Section 6. The diverse approaches are included in the summary both to document the scope of the available information, and to list alternatives which may be considered for future use.

The motions of an ice mass are governed primarily (1) by currents and winds, (2) by waves, and (3) by the hydrodynamic characteristics of the ice mass itself. The presence of a structure can affect all these factors. The ice-structure interaction after the initial instant of impact, depends also on the material properties of the ice mass but is not treated in this study. A brief but general review of this overall problem is given by Isaacson (1985).

The present study requires an understanding of the following related or component problems:

1. ice mass drift in open water and in the vicinity of a

large offshore structure,

2. wave-induced motions of a freely floating body (ice mass) in open water and in the vicinity of a large offshore structure.

The literature and data search has included a computer data base search using the following data bases: Compendex, NTIS, Oceanic Geoarchive, Engineering Meetings, and Fluidex. Keywords which were used in the searches have included added mass, hydrodynamic mass, hydrodynamic coefficients, iceberg, ice mass, floating body, barge, motion, drift, and wave. In addition, contact was made with several oil companies including Chevron, Mobil and Petro Canada; however, the review generally excludes reference to commercial projects which have not been reported in the open literature.

4.2 Ice Mass Drift

4.2.1 Drift In Open Water

Ice mass drift in open water is a relatively complex motion indicated by iceberg trajectories observed in the field. motions are predominantly rectilinear. typical observed circular and cusp-like trajectories are indicative of the complex and variable set of forces acting on the ice mass. Models proposed to describe this motion can broadly be classified as kinematic, dynamic, or statistical in nature depending on the approach to the problem. Kinematic models utilize a correlation ice mass velocities and known wind and velocities. Dynamic models predict trajectories by estimating the forces on the ice mass and integrating the equation of motion. Statistical methods base predictions on the past trajectory of the iceberg.

Kinematic models were amongst the earliest proposed. Their methodology ignores the actual forces on the ice mass and concentrates on the physical phenomena which give rise to the forces. Invariably, a linear relationship is assumed between the ice mass velocity and the velocities of the environmental factors in the form:

$$V_{i} = \sum_{n=0}^{\infty} k_{n} V \tag{4.1}$$

where V is the ice mass velocity, the V's are the velocities of the environmental factors, usually winds and currents, and the k's are the relative influence of each environmental factor. Dempster (1974) used a two term expression, one term representing a constant translational velocity and the second term representing a rotating tidal current. Cheema and Ahuja (1978) also used a two term representation; however in this case, ocean currents and winds were the environmental factors.

Dynamic models are more firmly based in the physical phenomena acting on the ice mass. The equation of motion can be written:

$$Ma = F + F + F + F$$
 (4.2)

where M is the iceberg mass combined with its added mass and is usually expressed in the form:

$$M = (1 + C_{M})M_{i}$$
 (4.3)

with M the iceberg mass; and C its added mass coefficient; a is acceleration; and F (a) M (c) and F are the forces of current drag, wind drag, Coriolis effect and waves respectively. The various dynamic models proposed differ mainly in the forces included and the value of the various coefficients required.

The current drag and the wind drag are computed from a quadratic relationship in the relative velocity as:

$$F^{(d)}$$
 or $F^{(a)} = 1/2\rho AC_{d} V_{r} |V_{r}|$ (4.4)

where ρ is the density of the fluid medium, A is the aerial or submerged cross sectional area normal to the relative velocity, C is the drag coefficient, and V is the velocity of the ice of mass relative to the wind or current. Wind is generally recognized as being of lesser importance and is often omitted from the model (Sodhi and Dempster (1975); Russel, Riggs, and Robe (1977)). Currents are a principal driving force and have been treated with various levels of sophistication. Smith and Banke (1981, 1982) utilized a uniform current while Sodhi and El-Tehan (1980) and Mountain (1980) utilized multi-layered models to separate surface, wind driven currents from mean currents.

Forces due to the Corolis effect result from the motion of the ice mass and the sloping sea surface which arises from the geostrophic current. These are combined into a single expression in the form:

$$F^{(C)} = M_{i}(f \times V_{r}) \tag{4.5}$$

where V is the velocity of the ice mass relative to the current and f is the component of the earth's angular velocity at latitude θ directed normal to the earth's surface with magnitude $2\Omega\sin\theta$ in which Ω is the earth's angular speed.

Wave forces are included in a model proposed by Hsiung and Aboul-Azm (1982). Only horizontal forces associated with long term drift were derived utilizing diffraction wave theory, ignoring oscillatory motions.

Various degrees of success have been claimed by the authors of these models. In general, the models predict trajectories which are predominantly rectilinear better than trajectories with a significant number of loops and cusps. A number of different drift models are regularly used in preparing iceberg position forecasts for East Coast offshore operations and are widely recognized as not sufficiently accurate for that application. The major limitation in drift prediction accuracy is the inaccuracies on the model input (wind, current, and wave data; iceberg shape and mass parameters; and drag and added mass coefficients).

Statistical models are designed to overcome the sensitivity of deterministic models to uncertainties and errors in Garrett (1984) has proposed a statistical scheme iceberg position and velocity data up to a given point in time to derive optimum linear estimators for future positions and velocities. Through the use of the Gauss-Markov coefficients leading to the lowest mean square deviation actual values are found. The iceberg's course is then projected future with a fading memory on the estimators. further the berg travels from its last recorded position, the greater the uncertainty in the estimates. The statistical model shows promise for predicting short term drift, up to 12 hours from the last data.

A hybrid model, developed by Gaskill and Rochester (1984), employs a dynamic analysis similar to that described above but estimates the input data from the previous history of a berg's drift.

4.2.2 Drift Near A Structure

Little work has been done on the two body problem in which an ice mass drifts within a few diameters of a structure. Obviously, the disturbances in the flow field caused by the ice mass and the structure lead to some interactions which will influence the drift. Isaacson (1985) has suggested an approach to the problem using boundary elements, a methodology often employed in diffraction wave theory and thus is amenable to superposition of

wave and currents effects on iceberg motions near a structure. The method involves expressing the velocity potential as a sum of components representing a uniform current and the disturbance created by the ice mass and structure. The potential for the uniform current is:

$$\phi_{C} = U(x\cos\beta + y\sin\beta) \qquad (4.6)$$

where U is the current velocity and β is its angle with the x axis of the co-ordinate system. The disturbance potential can then be derived using methods described in Section 4.3.

4.2.3 Data Input To Drift Models

The accuracy of trajectory projections depended on the quality of the input data. Three categories of inputs are required in most models:

- a. current, wind and wave conditions as functions of both space and time
- b. appropriate force coefficients
- c. iceberg parameters such as aerial and submerged shape, mass, orientation, and added mass.

Ball, Gaskill, and Lopez (1981) have assessed the problem by analyzing 250 iceberg trajectories observed in the field. Laboratory experimentation is also being utilized to refine drag coefficients (Shirasawa, Riggs, and Muggeridge (1984)).

A recent study, Arctic Sciences (1984), investigated the sensitivity of iceberg trajectory predictions to uncertainties in the inputs. The investigators found that the largest source of

uncertainty is associated with the water current information and the water drag force coefficient. They conclude that improvements required to obtain reliable predictions over large areas of the ocean are not likely in the near future.

Recognition that accurate long term drift predictions will not soon be available has focussed attention on short term drift in the vicinity of offshore structures. Lever (1984) and Arctec (1984), have performed laboratory investigations of wave and current induced drift which ultimately will be helpful in the development of a short term drift model.

4.3 Wave-Induced Motions Of Floating Bodies

4.3.1 Flow Regimes

Wave-structure interaction problems generally encompass two distinct flow regimes, one associated with small bodies, which behaviour is governed by flow separation, (drag-dominated), other with large bodies, which diffract the (inertia-dominated). For small diameter fixed structure, flow separation occurs and influences the loading on structure, and the waves propagate essentially unaffected. Α large diameter structure diffracts the incident wave field and the resulting pressure field determines the load on the structure.

In the case of a freely floating body such as an ice mass, there are three distinct regimes. Sufficiently small ice masses essentially behave as fluid particles and follow the orbital motion of the fluid. For larger ice masses wave diffraction is the predominant factor in determining the motions of the body. Typically, diffraction becomes a significant influence when (diameter/wave length) exceeds about 0.2. For extremely large ice masses: (typically D/L > 2.0) the body still diffracts the

waves but its oscillatory motions are negligible. The presence of a large structure in the vicinity of the ice mass modifies the wave field and produces a reflected wave field immediately upstream of the structure.

4.3.2 General Calculation Method

of wave loading and structural response problems Reviews involving fixed and/or floating structures have been given several texts and general reviews (e.g. Wehausen (1971), (1974), Hogben et al. (1977), Newman (1977), Garrison (1978), Mei (1978, 1983), Isaacson (1979), Sarpkaya and Isaacson (1981), Yeung (1982), Hooft (1983), etc.). Most computer models of motion response in waves are based on linear wave theory in which the body oscillates sinusoidally in a regular wave train of small amplitude. The body may oscillate with six degrees of freedom corresponding to surge, sway, heave, roll, pitch, and yaw, denoted by ζ , j=1,...,6. In a linear analysis, the fluid force acting on the structure may be separated into components proportional to the acceleration and velocity components of the body motion, expressed as added-mass and damping coefficients respectively, and to a remaining component which involves incident waves, termed the exciting force component. equations of motion of the body may therefore be expressed as:

$$\{-\omega^2([M] + [A]) - i\omega[B] + [C]\}(\zeta) = (F_e)$$
 (4.7)

Here [M] is the mass matrix of the body, [C] is the stiffness matrix of the body associated with hydrostatic stiffness components, [A] is the added mass matrix, [B] is the damping coefficient matrix, and (F) is the exciting force vector. In general, [M] and [C] can be estimated from the configuration of the ice mass. The matricies [A] and [B] and the vector (F) may be obtained by a hydrodynamic analysis. For a small structure, or one composed of slender structural elements, this involves the

use of the Morison equation, with a suitable linearization of the drag force. For a large structure, this involves a solution to the related diffraction and radiation problems. Once [A], [B] and (F) are all known, Eq. 4.7 may be solved to obtain the motions (ζ) .

The analysis for motions in regular waves may be extended to the case of irregular waves by means of the superposition principle The method involves solving Eq. 4.7 to linear theory. determine the response amplitude operators, RAO's, for a series regular incident waves of unit amplitude and spanning a range of frequencies. If required, the RAO's can then be applied to obtain time histories of the body response to a irregular wave record. More commonly, they can be used conjunction with a specified wave spectrum $S_n(f)$ to obtain the corresponding spectra $S_r(f)$ of the component motions:

$$S_{\zeta}(f) = \{RAO(f)\}^{2}S_{\eta}(f)$$
 (4.8)

These response spectra can then be integrated to provide the root-mean-square motions and in turn the maximum expected motions in a sea state of specified duration. If required, a simple extension to this analysis provides the accelerations at various specified points on the body.

4.3.3 Small Floating Bodies

The analysis procedure for a floating body which is small relative to the wave length (or one which is composed of small diameter structural members) is treated in several reviews, including the texts of Sarpkaya and Isaacson (1981) and Hooft (1983). In such a case the Morison equation is used to express the force on a segment of each member in terms of the incident flow and velocities and accelerations of the structure. In order to develop the equations of motion into the form of Eq. 4.7, the

nonlinear drag term must be approximated by a linear term.

If the body consists of a single member (corresponding to an ice mass which is small relative to the incident wave length) then it generally follows the water particle motions providing no resonance occurs. A body with acceleration $\dot{\mathbf{u}}_b$ immersed in a fluid which itself has an acceleration $\dot{\mathbf{u}}_f$ experiences a force F equal to (ignoring resonance):

$$F = \rho V(1 + k)\dot{u}_{f} - \rho Vk\dot{u}_{b}$$
 (4.9)

where k is an added-mass coefficient, is the fluid density and V is the immersed body volume. The equation of motion of the body whose mass is ρV is $F = V\dot{u}$ and thus $\dot{u} = \dot{u}$. Since the relative velocity $\dot{u} - \dot{u}$ is small, the drag force is insignificant and the body follows the fluid motion.

4.3.4 Large Floating Bodies

For a large body which diffracts the incident waves, diffraction theory may be used to calculate the wave-structure interaction. This topic has been reviewed by a number of authors including Hogben (1974), Garrison (1978), Mei (1978, 1983), Isaacson (1979a), Sarpkaya and Isaacson (1981) and Yeung (1982). A number of theoretical and numerical methods are available for treating such problems. Some treat general three-dimensional bodies of arbitrary shape, others treat bodies with a two-dimensional irregularity and others are restricted to specific reference configurations.

(a) General Three-dimensional Bodies

The general case of three-dimensional structures of arbitrary shape may be treated numerically by either source distribution methods or finite element methods. In the former methods the surface of the submerged structure is discretized into a number

of surface elements or facets, [Garrison and Chow (1972), Faltinsen and Michelsen (1974), Hogben and Standing (1974), Loken and Olsen (1976), etc.] In the finite element methods the fluid domain (within a suitable exterior boundary) is discretized into three-dimensional elements [Mei (1978), Zienkiewicz, Bettess and Kelly (1978), etc.].

For ship-like structures which have a well-defined longitudinal axis, methods based on strip theory may also be used. The structure is divided into a series of parallel vertical planes and thereby extending two-dimensional methods to three-dimensional structures. This extension has been summarized by Newman (1977).

(b) Two-dimensional Configurations

Many structural configurations possess a two-dimensional rather than a three-dimensional irregularity, and only a discretization of the structure's profile (for boundary element methods) or of the two-dimensional fluid domain (for finite element methods) is needed. Three fundamental configurations possessing such a two-dimensional irregularity are:

- a. vertical axisymmetric bodies,
- b. horizontal cylinders of arbitrary section, and
- c. vertical cylinders of arbitrary section.

The wave force on a fixed vertical axisymmetric body has been considered by several authors, including Fenton (1978). The extension to floating axisymmetric bodies, which includes the numerical calculation of hydrodynamic coefficients and response amplitude operators, has been treated by Isaacson (1982c). A boundary element method is used, with the body's axisymmetry utilized to reduce considerably the computational effort required from that for bodies of arbitrary shape.

The second case above corresponds to a two-dimensional problem in the vertical plane and a two-dimensional boundary element method employing a suitable Green's function or the finite element method may be employed to treat this. Corresponding programs for floating bodies have been described by a number of authors, including Adee and Martin (1974) and Maeda (1974) using the boundary element method and Bai and Yeung (1974) and Isaacson and Fraser (1979) using the finite element method. The extension to treat ship shapes by strip theory is outlined by Newman (1977).

The third case has been treated by Isaacson (1978a) for fixed structures, but this has not been extended to floating structures.

(c) Reference Configurations

A variety of specific solutions have been obtained for particular reference configurations. For example, various kinds of vertical circular cylindrical bodies that have been analyzed include a single surface-piercing cylinder, MacCamy and Fuchs (1954), more neighbouring surface-piercing cylinders, Spring Monkmeyer (1974), Chakrabarti (1978), McIver and Evans (1984), a truncated cylinder at the free surface, Black, Mei and (1971), Garrett (1971), a truncated cylinder resting seabed. Black, Mei and Bray (1971), an axisymmetric compound cylinder with components of unequal radii, Isaacson (1979b) so on. These different solutions are usually very specific so as not to be directly applicable to the study problem.

4.3.5 Nonlinear Wave Effects

The full Morison equation with appropriate force coefficients is required with a suitable nonlinear wave theory. However, the extension to force spectra associated with irregular waves is more difficult and the problem is usually analyzed in the time

domain.

The extension for the case of large bodies is much more involved. Two separate approaches have been developed to treat the problem. based on a perturbation procedure, in which linear diffraction theory is extended to a second order on the basis expansion procedure. The fundamental case of surface-piercing fixed vertical circular cylinder has received particular attention (e.g. Rahman (1984) and a number of papers referred to therein). The more general case structure of arbitrary shape has been treated by Lighthill (1979) who describes the general formulation of the problem (for a fixed structure) but does not obtain any results, and also by Garrison The solution involves a second order integration over the free surface in the vicinity of the structure, thus increasing the degree of computational effort in relation to the first order problem.

A second approach has been to solve the complete nonlinear equations numerically, which requires a time stepping procedure since the surface boundary conditions must be applied at the instantaneous free surface which is continuously changing. The methods which have been developed are all essentially boundary integral methods, in which a boundary integral equation is solved at each time step to provide the required variables over the fluid boundary.

Faltinsen (1977) treated the two-dimensional vertical plane case of an oscillating cylinder in deep water. More recently, Vinje and Brevig (1981) and Vinje, Xie and Brevig (1982) have used a similar procedure to describe the development of wave breaking and wave interaction with a fixed or oscillating cylinder, again for the two-dimensional vertical plane case. Stansby and Slaouti (1984) have treated the vertical plane problem using a vortex sheet approach.

A numerical treatment of the complete three-dimensional nonlinear diffraction problem involving a floating structure has been given by Isaacson (1982a, b). For example, the heave and surge response of a floating circular cylinder subjected to a specified incident wave condition are presented and the extent of the nonlinear effect is demonstrated.

4.3.6 Interference Effects

The interference effect between large neighbouring structures in waves has application to the behaviour of an ice mass in the vicinity of a large offshore structure. This topic has been studied in the past to assess the interference effects for such situations as neighbouring offshore storage tanks, neighbouring columns of a gravity platform, neighbouring hulls of a vessel, etc.

The interference effect between fixed vertical circular cylinders been studied by several authors including Ohkusu (1974), Spring and Monkmeyer (1974), Chakrabarti (1978), Isaacson (1978b) and McIver and Evans (1984). Generally, the interference is most severe when the cylinders are aligned with the wave direction and is associated with the upstream cylinder lying in the reflected wave system of the other. Isaacson (1978) reports that even for cylinders which are ten radii centre to centre apart, the line force on one cylinder may be increased by as much as When cylinders are aligned with the incident wave crests interference effect is relatively small. Matsui and an extension to this problem dealing with interference effect between groups of vertical axisymmetric in waves and presented results for a pair of circular cylinders of finite draft.

Van Oortmerseen (1979, 1981) considered the interference effect between neighbouring floating cylinders, with one cylinder circular and the other almost square and both with the same draft. In general he found that the presence of a neighbouring structure results in added mass and damping coefficients versus frequency curves that oscillate around the corresponding curves for an isolated structure, and that the interference effect on these diminish as the frequency tends to zero. A structure is also subject to forces resulting from motions of the neighbour and results of these interaction coefficients were present over the entire frequency range. Loken (1981) subsequently carried out a similar study and compared his results with those of van Oortmerseen (1979).

The case of neighbouring horizontal cylinders has also been considered by several authors, mainly in the context of interference effects between neighbouring hulls of a vessel, and corresponding to a two-dimensional vertical plane problem. This has been treated by several authors including Wang and Wahab (1971) and Ohkusu (1974), who have considered pairs of semi-immersed cylinders of various sectional shapes and have presented results on exciting forces, added-mass and damping coefficients and response amplitude operators.

The interference between neighbouring bodies immersed in a uniformly accelerating flow of an infinite fluid has been considered by Yamamoto (1976). He provides expressions and results for the hydrodynamic forces on an arbitrary number of cylinders in arbitrary motion based on potential theory and considers several special cases. The force exerted on a moving cylinder in the presence of a fixed neighbour and an otherwise stationary fluid is of interest here and is given as:

$$F = - \pi \rho a^{2} C_{1b}^{\dot{u}} + \rho a C_{3} | u_{b} |^{2}$$
 (4.10)

where a is the cylinder radius and the coefficient C is the added mass coefficient. (C = 1 for an isolated cylinder). C is a convective force coefficient which arises from the velocity

squared terms in the Bernoulli equation for pressure. (C = 0 for an isolated cylinder). As an example of the results obtained, for two cylinders of equal radii, C increases by about 10% when the gap width e = a, and by about 25% when e = 0.2a. The other coefficient C is about 0.1 when e = a and increases to about 1.0 when e = 0.2a. This force always acts in a direction so as to attract the two bodies, and thus opposes the drag force when one cylinder is approaching the other.

4.3.7 Viscous Damping of Large Bodies

Additional damping associated with viscous effects may have a significant influence on resonant oscillations of a large floating body. This may arise from shear stresses associated with the boundary layer on a smooth structure surface such as a ship's hull, or from vortex shedding off the sharp edges of a body. The latter case is important in the prediction of roll damping and resonant roll amplitudes for rectangular-section barges or other sharp-edged floating vessels. Graham (1980) has carried out a fundamental investigation into the forces associated with vortex shedding off a sharp edge in oscillatory flow and provides estimates of the forces which arise.

For roll motions of a vessel, the viscous damping force is nonlinear and dependent on the roll amplitude, so that the complete solution for roll response requires a time integration method. Alternatively, it may be convenient to adopt a method of equivalent linearization, with the equivalent linear damping taken as a function of the roll response. These procedures have been reviewed by Kaplan (1983).

4.3.8 Low Frequency Hydrodynamic Coefficients

A significant number of studies deal with the added mass and damping coefficients of floating bodies in waves. The low frequency limit is of particular interest to the study problem since this affects the drifting motion of an ice mass, and this topic has received some attention in the context of low frequency drifting oscillations of moored vessels. The damping coefficients (due to wave damping only) approach zero in the limit of zero frequency, but the added-mass coefficients do not.

examples of experimental and numerical studies, Wichers van Suijj (1979) provided data on damping coefficients associated with low frequency motions of tankers, and subsequently Wichers low frequency discussed the damping coefficients associated with radiation and viscous damping for tankers compared measured and computed damping coefficients and drift model tankers. Chakrabarti (1984)reported motions for experimental measurements of added-mass and damping coefficients for several structures which included a barge, tanker, mersible and a floating vertical cylinder. The dependence of these on wave frequency was presented.

4.3.9 Response of Ice Masses

Relatively few studies address wave effects on ice masses themselves. Wave effects on larger icebergs and ice floes have been studied mainly in the context of iceberg motions and breakup due to wave-induced flexural failure. Schwerdtfeger (1980) carried out a simple linear analysis to determine the heave and pitch resonance frequencies of a berg, but did not predict its motion amplitudes in waves. Orheim et al.(1982) reported on measurements of surface strain and heave motions of large tabular icebergs in waves. Squire (1981) and Kristensen and Squire

(1983) reported on a field study and theoretical analysis of the strains and motions of large ice in waves. Wadhams, Kristensen and Orheim (1983) measured the heave and roll motions and strains for three Antarctic icebergs and correlated these to wave measurements.

With regard to smaller ice masses, Lever et al. (1984) describes a laboratory study of the wave-induced motions of small ice masses corresponding to prototype masses in the range 10 to 10 tonnes. They concluded that for D/L less than 1/13, where D is the ice mass diameter, the ice masses essentially act as fluid particles and follow the wave flow. When heave resonance occurs, however, amplification factors up to 3.6 were measured. For D/L greater than about 1/10 the ice mass motion differed from the particle motion, due to viscous and/or diffraction effects.

5. SURVEY OF AVAILABLE COMPUTER MODELS

Due to the wide range of environmental conditions and possible ice mass and structure sizes and shapes, a prime objective of this study is the acquisition and/or development of an appropriate computer model.

During the past 15 years, considerable model development has been support of offshore oil activities. Models predicting the response of floating bodies to waves and wave on structures have been developed and are available commercially from many sources. In order not to duplicate computer models which might provide a previous work, starting point for the study objectives here were investigated.

Oil drilling activity off Canada's East Coast has also led to the development of computer models predicting the drift of icebergs open sea. As described in Section 4.2, the forces on an iceberg due to wind, compute currents, and include forces that are important in long term drift predictions such as the Coriolis force. None of these models, however, has the option of including a second, fixed body. their features are superfluous to ice motions structure and since their formulations ďο not permit modifications to include a structure, models were these considered for evaluation.

The unique demand of the present study, relative to the general capabilities of available models, is the inclusion of the second, fixed body. Models whose principal component is a three dimensional diffraction and radiation analysis (Section 4.3) hold the best hope for significant capabilities in the general area of interest. Numerous such computer models developed by companies undertaking research and development or design in support of shipbuilding or offshore oil drilling were identified. Many are proprietary to the companies which developed them and some are

available for purchase under limiting licensing arrangements.

Most of the models were originally developed for specific applications, and cannot be readily adapted to different types of For example, one candidate program computes currentinduced drift, but cannot accommodate distortions to the current field caused by a fixed or floating body. A number of programs were identified which address multi-body problems. assume a complete de-coupling of the these, however, interactions from the wave and current fields. acceptable for the originally simplifications are applications, but not for the present study. To allow a solution to the two-body problem, de-coupling assumptions are also introduced (Section 6.1.3) in this study. These assumptions, however, are far less restrictive than those used in past work, and retain the key features of the interactions.

Major practical restrictions, which could not be addressed within the resources of this study, precluded the use of most commercially-available models, requiring modifications for the study problem. These limitations include:

- (1) some of the models are modules of larger suites of programs, and are not designed to be easily extracted and developed separately;
- (2) under licensing agreements, some of the models can not be modified by anyone other than the vendor and thus do not meet the terms of reference for the study;
- (3) even within licensing arrangements, the task of becoming familiar with the logic and programming of the model would take more time than was available to the study;
- (4) the cost of purchasing the models was well beyond the

means of the study.

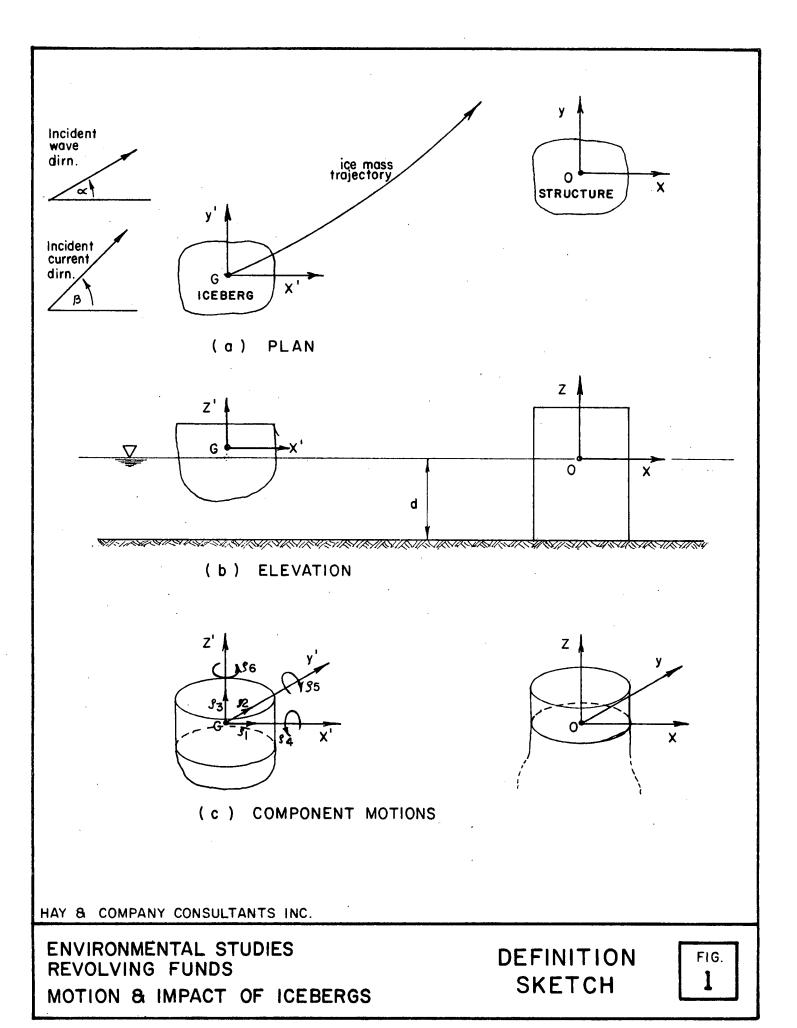
One model which did not suffer from the above constraints is that developed by Prof. M. Isaacson who is a principal investigator in this study. His model was developed as a research tool and included a single-body diffraction and radiation program. Currents were not included in the model and, therefore, it was not as advanced in that sense as some of the commercially available models. Owing to the problems faced by the commercial models though, the Isaacson model was adopted as a starting point.

6.0 THEORETICAL AND NUMERICAL DEVELOPMENT OF THE MODEL

6.1 Background

This section describes the theoretical and numerical development of computer programs for treating motions of an ice mass due to waves and currents in close proximity to a large fixed offshore The general situation is indicated in Figure 1. structure. fundamental approach superposes the wave-induced oscillatory of the ice mass on the drift motion due to the that two component problems are treated. induced motions, obtained by a diffraction calculation, depend on ice mass proximity to the structure. The motions must be recalculated at a series of points along the trajectory. Furthermore, the drift may be influenced by wave drift forces and by the zero frequency added mass components the ice. Both of these terms also vary with the location of the ice mass and may be obtained by the wave diffraction calculation. Thus the two component problems, current drift and wave-induced motions, are coupled.

development below is similar in many respects to those well-established single-body problems. Though formulations single-body problem are available in standard references Sarpkaya and Isaacson, (1981) a complete, self-contained development is provided here for two reasons. First, computer model can be prepared using this section as a basis; and secondly, existing single-body models can be extended to two-body models, also using this section. As part of this study, a twobody computer model was developed from a single-body program. original program (Section 5) is proprietary, a significant portion of the coding of the final program is also. Therefore, inclusion of the full theoretical and numerical developments circumvents the difficulties of proprietary coding.



6.1.1 Wave-Induced Motions

Three distinct flow regimes are associated with the wave-induced For ice masses of a diameter to wave motions. length D_i/L between approximately 0.2 and 2, the oscillatory motions of mass are not negligible. Wave diffraction and radiation by ice mass as well as wave diffraction by the structure significant and must be considered in the ice mass motions. These are obtained by solving the two body linear diffraction problem in which one body (the structure) is fixed and the other (the ice mass) is freely floating. For very large ice masses, corresponding to D_i/L greater than about 2, both the ice mass and the structure diffract the waves but the ice mass oscillatory motions are negligible. The wave drift force, is required for the drift motion computation so that a diffraction computation is necessary. For smaller ice corresponding to D_i/L less than about 0.2, wave diffraction from the ice mass is negligible. The ice mass then essentially behaves a fluid particle and follows the orbital motion of the fluid. In this case, only a single, fixed body diffraction computation required to obtain the orbital velocities at the location. For certain ranges of parameters, however, oscillating motions may be increased by resonance effects heave, pitch and roll. The calculation of these effects requires calculations of the added masses and damping characteristics the ice mass which are obtained from the two body diffraction solution.

Overall, a two body diffraction calculation is usually necessary over the full range of ice mass sizes, although a single fixed body diffraction calculation may sometimes be adequate.

In the calculation of wave-induced motions, the ice mass is considered to possess all six degrees of freedom, corresponding to surge, sway, heave, roll, pitch and yaw.

6.1.2 Drift Motion

Different flow regimes also apply for drift motions. When the ice mass diameter to structure diameter D /D is sufficiently small, the ice mass may be considered to be located in the nonuniform current field as influenced by the structure. The trajectory of the ice mass then obtained by a time-stepping procedure applied to this constant current field. If D /D is sufficiently large, the extent of the disturbed flow field around the structure is on the same order as the size of the ice mass and, therefore, has a smaller influence on the ice mass. In this case, a uniform current field may be assumed and the presence of the structure on the current field ignored.

With respect to drift motions, the ice mass possesses two degrees of freedom - surge and sway.

The drift motions may depend also on the wave drift forces and zero frequency added masses. The relative importance of wave drift forces may be assessed with reference to the ratio of wave drift force F to current drag force F, for a floating truncated circular cylinder. This is given approximately as:

$$\frac{\mathbf{F}^{(\mathbf{w})}}{\mathbf{F}^{(\mathbf{d})}} = 0 \left[\frac{0.1 \text{gH}^2}{\frac{2}{\text{U}_{\mathbf{r}}^2}} \right]$$
 (6.1)

where g is the gravitational constant, H is the wave height, U is the current magnitude relative to the drift velocity of the ice mass and h is the draft of the ice mass. Thus for any specified situation, this ratio may be used to assess the need to include wave drift forces in the drift computation.

The zero frequency added masses also appear in the drift

equations of motion and depend on the proximity of the ice mass to the structure. They are obtained by solving the two body diffraction problem at a very low wave frequency. Certain simplifications may be applicable in this calculation, particularly in the Green's function used and in the absence of any incident flow.

6.1.3 Underlying Assumptions

A complete modelling of the full physical problem including the numerous nonlinear and secondary influences, is not at present possible. Assumptions, valid over a practical range of physical conditions and parameters, are introduced to render the problem solvable. The chief implicit assumptions are:

- 1) The current is sufficiently slow so that free surface effects associated with the current field in the absence of waves can be neglected.
- Nonlinear interaction terms are sufficiently small so that the wave and current fields are each treated separately.
- In the wave diffraction calculations, the ice mass is assumed to have a fixed equilibrium position (i.e., the oscillatory motions and wave radiation are not influenced by the superposed mean drift). This quasisteady assumption is valid for relatively slow drift velocities. Typically, the fluid particle velocities due to waves are an order of magnitude greater than those due to currents.
- 4) Likewise, the drift forces are calculated using the mean, rather than the total (mean plus oscillatory) ice mass velocities. As the wave-induced motions are zero-

mean and as the wave drift is calculated separately, the assumption is appropriate. Assumptions 1 through 4 describe the de-coupling of wave and current interactions introduction by the proposed solution.

- 5) As is commonly adopted in diffraction analyses, the wave height is assumed small relative to the wavelength.
- 6) The fluid is incompressible and inviscid. The flow is assumed irrotational to allow a solution in terms of velocity potential. Flow separation is therefore initially neglected. Viscous damping effects on the ice motion are included through the use of empirical coefficients in the equations of motion.
- 7) Yaw motion in the drift is ignored. The assumption is independent in that it is a consequence of the treatment of the drift flow regimes described in Section 6.1.2.
- 8) In Section 6.1.2, the effect due to wind, Coriolis force and sea-surface slope are ignored.

6.1.4 Computer Procedure

A computer procedure, based on the above considerations, is described by the flow chart given in Figure 2. In essence, initial diffraction and current field calculations are first carried out, and then a time-stepping procedure is applied to the drift equations of motion in order to obtain successive values of the drift velocity and location of the ice mass. At each time step the current field at the ice mass location is recalculated, and, after a user-specified number of time steps, the diffraction calculation is repeated.

6.2 Wave Induced Motions

The method described below is based on a boundary element method utilizing a three-dimensional Green's function. It is applicable to a two-body problem, with one fixed and the other floating in regular waves. The notation and conventions consistent with those used in common formulations of the singlebody problem (Sarpkaya & Isaacson, (1981)). The extension to random waves is straightforward in principle and essentially requires a repetition of the diffraction calculation for a series wave frequencies. The extension requires a large number separate diffraction computations to be carried out because calculation must also be performed for a series of ice mass locations. The computations would thus be extremely and were not attempted in this study. A further extension directional seas is possible in a similar way, but again would substantially increase the computational effort required.

6.2.1 Governing Equations

A regular small amplitude wave train propagates in water constant depth d past a fixed offshore structure and freely floating ice mass, both of arbitrary three-dimensional shape Figure 1. Since the ice mass is drifting and its component motions are conveniently described with respect to axes fixed relative to its drifting location, it is appropriate employ two coordinate systems as indicated on the figure. is the centre of gravity of the ice mass and Gx'y'z' forms a right-Cartesian coordinate system relative to which the motions of the ice mass are defined and with axis vertical. Oxyz forms a second Cartesian coordinate system, fixed in space, with the origin 0 conveniently located at a reference point within the structure and at the still water level. The axis x is measured horizontally and parallel to the axis Gx' of the ice mass and the axis z measured upwards.

Although the ice mass is drifting, the wave-induced motions are treated by considering the ice mass to have a fixed equilibrium position so that for any one computation the Gx'y'z' coordinate system is considered fixed. Three parameters are then needed to determine the relationship between the two, parallel coordinate systems. This conveniently involves x_{G} , y_{G} , z_{G} , the x, y, z coordinates of G, and is simply:

$$x' = x - x_G$$

$$y' = y - y_G$$

$$z' = z - z_G$$
(6.2)

The wave train has a height H and angular frequency ω and propagates in a direction α to the x axis. The fluid is assumed incompressible and inviscid, and the flow irrotational so that the fluid motion may be described by a velocity potential (defined here such that the fluid velocity vector $\underline{\mathbf{u}} = \nabla \phi$) which satisfies the Laplace equation within the fluid region. The wave height is assumed sufficiently small for linear wave theory to apply so that ϕ is subject to the usual boundary conditions, linearized where appropriate, at the seabed; the ice mass and structure submerged surfaces; the free surface; and the far field.

In general, the ice mass oscillates with six degrees of freedom corresponding to surge, sway, heave, roll, pitch and yaw, denoted by the subscripts 1, ..., 6 respectively, and defined with respect to the Gx'y'z'system as indicated in Figure 1. Each mode of motion is harmonic and may be expressed in the form ζ_k exp $(-i_{\omega} t)$, in which t is time and ζ_k is the complex amplitude of each component motion.

Since the drift motion over a few wave periods is small, the

linearized boundary condition on the ice mass surface is applied on its equilibrium surface S which is considered fixed. Thus this boundary condition on the submerged equilibrium surface of the ice mass S and on the submerged surface of the structure S is given as:

$$\frac{\partial \phi}{\partial n} (\underline{x}) = \begin{bmatrix} 0 & \underline{x} \text{ on } S \\ & & \underline{x} \text{ on } S \\ & & & \underline{x} \text{ on } S \end{bmatrix}$$

$$(6.3)$$

where \underline{x} represents a point in the fluid, and n denotes distance in the direction of the unit vector \underline{n} directed normally out of S. Also V is the velocity of the ice mass surface in the direction \underline{n} and is itself given as:

$$V_{n}(\underline{x}) = \sum_{k=1}^{6} -i\omega \zeta_{k} c_{k}(\underline{x}) \exp(-i\omega t)$$
 (6.4)

in which

where n , n , n are the direction cosines of \underline{n} in the x,y,z (or x',y',z') direction respectively.

The flow potential is made up of components associated with the incident waves (subscript 0), the scattered waves due to both the

structure and ice mass (subscript 7) and forced waves due to each mode of motion (subscripts 1, ..., 6), each of the latter being proportional to the corresponding motion amplitude. Thus, it is convenient to express the total flow potential as:

$$\phi = \left\{-\frac{i H}{2} \left(\phi_0 + \phi_7\right) + \sum_{k=1}^{6} -i\omega\zeta_k\phi_k\right\} \exp(-i\omega t)$$

$$(6.6)$$

in which $\boldsymbol{\varphi}_k$ is generally complex.

Substituting Eqs. 6.4-6.6 into Eq. 6.3, and separating out terms corresponding to the diffraction problem (k = 0, 7) and to each component of the radiation problem (k = 1, ..., 6), the body surface boundary condition can be decomposed into the form:

$$\frac{\partial \phi}{\partial n} k(\underline{x}) = \frac{1}{2} b_k(\underline{x}) \tag{6.7}$$

where:

$$b_{k}(\underline{x}) = \begin{bmatrix} 0 & k=1,...,6 \text{ and } \underline{x} \text{ on } S \\ 2n_{k}(\underline{x}) & k=1,...,6 \text{ and } \underline{x} \text{ on } S \\ & & \vdots \\ -2\frac{\partial \phi}{\partial n} (\underline{x}) & k=7 \text{ and } \underline{x} \text{ on } S \end{bmatrix}$$

and S represents the combined surfaces S $_{\rm i}$ and S $_{\rm s}$

The incident potential ϕ is known and is given as:

$$\phi_0 = \frac{\cosh(k(z+d))}{k\sinh(kd)} \exp[ik(x\cos\alpha + y\sin\alpha)]$$
 (6.9)

where k is the wave number.

6.2.2 Source Distribution Representation

The unknown potentials may each be represented by a distribution of point wave sources over the surface S. Thus the potential $\phi_k(\underline{x})$ at a general point \underline{x} within the fluid is expressed as:

$$\phi_{\mathbf{k}}(\underline{\mathbf{x}}) = \frac{1}{4\pi} \int_{S} f_{\mathbf{k}} (\underline{\xi}) G(\underline{\mathbf{x}}, \underline{\xi}) dS \qquad k=1,...,7$$
 (6.10)

where $f_k(\xi)$ represents a source strength distribution function, ξ is a point on the surface S over which the integration is performed, and $G(\underline{x}, \underline{\xi})$ is a Green's function for the general point \underline{x} due to a source of unit strength located at ξ . G is itself chosen to satisfy the Laplace equation, the seabed boundary condition, the linearized free surface condition and the radiation condition. Alternative expressions for G have been developed. (See, e.g. Hogben & Standing, 1974).

The source strength functions f may be determined by applying the body surface boundary condition. This condition, given by Eqs. 6.7 and 6.8, together with the representation for ϕ given by Eq. 6.10, gives rise to a set of surface integral equations for f:

$$-f_k(\underline{x}) + \frac{1}{2\pi} \int_S f_k(\underline{\xi}) \frac{\partial G}{\partial n}(\underline{x}, \underline{\xi}) dS = b_k(\underline{x})$$

$$k = 1, ..., 7 (6.11)$$

where \underline{x} now lies on the surface S at the point where the boundary condition is applied, n is measured from \underline{x} , and the integration is carried out over the points ξ .

6.2.3 Discretization Procedures

In a numerical solution to Eq. 6.11, the surface S is discretized into a number of small facets, with a total of N facets being made up of N and N facets on S and S respectively, and the functions f are taken as uniform over each facet. The integral equations are then applied at each facet centre in turn and so may be expressed as seven sets of linear algebraic equations:

$$\sum_{j=1}^{N} A_{ij} f_{j}^{(k)} = b_{i}^{(k)} \qquad i=1,...,N \quad k=1,...,7$$
(6.12)

where f (k) and b (k) denote f (x) and b (x) respectively, and x is the value of x at the centre of the j-th facet. If the facet sizes are small relative to the wave length, an approximate expression for the coefficients A may be developed as follows:

$$A_{ij} = \begin{bmatrix} \frac{\Delta S}{2\pi} j & \frac{\partial G}{\partial n}_{i} & (\underline{x}_{i}, \underline{\xi}_{j}) & & i \neq j \\ -1 & & i = j \end{bmatrix}$$
(6.13)

where Δ S is the area of the j-th facet. Equations 6.11 may now be solved to determine the source strengths f . Once the source strengths are known, the potentials ϕ at the facet centres on the ice mass may themselves be obtained from a discretized form of Eq. 6.10:

$$\phi_{i}^{(k)} = \sum_{j=1}^{N} B_{ij} f_{j}^{(k)}$$
(6.14)

where ϕ denotes $\phi_k(\underline{x})$. An approximate expression for the coefficients B may be obtained in a similar way as for A and is:

$$\mathbf{B}_{\mathbf{i}\mathbf{j}} = \begin{bmatrix} \frac{\Delta \mathbf{S}}{4\pi} \mathbf{j} & \mathbf{G}(\mathbf{x}_{\mathbf{i}}, \underline{\xi}_{\mathbf{j}}) & & & & \\ \frac{\Delta \mathbf{S}}{4\pi} \mathbf{i} & \Gamma & \Delta \mathbf{S}_{\mathbf{i}} & \mathbf{G}(\mathbf{x}_{\mathbf{i}}, \underline{\xi}_{\mathbf{i}}) & & & & \\ & & & & & \\ \end{bmatrix}$$
 (6.15)

where

$$\Gamma = \frac{1}{\sqrt{\Delta S_i}} \int_{\Delta S_i} \frac{1}{R} dS$$
 (6.16)

 $R = |\underline{x} - \underline{\xi}|$ and $G(\underline{x}, \underline{\xi})$ excludes the singular term 1/R which appears in the expression for G. The integral Γ may be obtained for any given facet shape as indicated by Hogben and Standing (1974).

6.2.4 Hydrodynamic Coefficients

The components of fluid force or moment acting on the ice mass are expressed as:

$$F_{j} = -i\omega \int_{S_{i}} \phi n \, dS \qquad j=1,\dots,6$$
(6.17)

Here F , F , F denote the force components along the x', y', z' axes and F , F , F denote the moment components about the x', y', z' axes as indicated for ζ in Figure 1. Each such component may be decomposed into an exciting force component F associated with ϕ + ϕ , and components associated with the forced potentials ϕ , ..., ϕ . The latter may in turn be described in terms of added masses μ and damping coefficients λ (which jk are real) so that F is expressed as:

$$\mathbf{F}_{\mathbf{j}} = \mathbf{F}_{\mathbf{k}=1}^{(e)} + \sum_{\mathbf{k}=1}^{6} (\omega^{2} \mu_{\mathbf{j}k} + i\omega\lambda_{\mathbf{j}k}) \zeta \exp(-i\omega t)$$
 (6.18)

Expressions for $\frac{F}{F}$, μ and λ may be obtained by substituting Eqs. 6.6 and 6.18 into Eq. 6.17 and collecting corresponding terms. Using the discretization approximation already described, this gives:

(e)
$$\sum_{j=1}^{N_{i}} (0) (7)$$
 $\sum_{j=1}^{N_{i}} (6.19)$

$$\mu_{jk} = -\rho \operatorname{Re}\{I_{jk}\}$$
 (6.20)

$$\lambda = -\rho\omega Im\{I\}$$
jk (6.21)

where

$$I_{jk} = \sum_{i=1}^{N_{i}} (k)$$

$$\sum_{i=1}^{N_{i}} \phi_{i} \quad n_{ji} \Delta S$$

$$(6.22)$$

n denotes n (x) and Re $\{\}$ and Im $\{\}$ denote the real and imaginary parts respectively.

6.2.5 Response Amplitude Operators

Once the various hydrodynamic coefficients described above have been obtained, it is relatively straightforward to solve the equations of motion of the ice mass in order to determine the six components of motion. The equations of motion may be written in the form:

$$\sum_{k=1}^{6} \{-\omega^{2}(m_{jk} + \mu_{jk}) - i\omega(\lambda_{jk} + \lambda_{jk}^{(v)})\} = c_{jk} \zeta_{k} = F_{j}^{(e)}$$
 (6.23)

where m is the mass matrix, c is the hydrostatic stiffness jk matrix and λ (v) are empirical viscous damping coefficients which are taken as zero except possibly for heave, roll and pitch motions.

The mass matrix components are given as:

$$\mathbf{m}_{jk} = \begin{bmatrix}
\mathbf{m} & 0 & 0 & 0 & 0 & 0 \\
0 & \mathbf{m} & 0 & 0 & 0 & 0 \\
0 & 0 & \mathbf{m} & 0 & 0 & 0 \\
0 & 0 & \mathbf{m} & 0 & 0 & 0 \\
0 & 0 & 0 & \mathbf{I} & -\mathbf{I} & -\mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
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0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
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0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
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0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
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0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
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0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} & \mathbf{I} \\
0 & 0 & 0 & -\mathbf{I} \\$$

where m is the mass of the ice and I , I , I are its moments of inertia about the x', y', z' axes respectively. Also, I , I and I are its products of inertia defined as: 12

$$I_{12} = \int_{V} \rho_{i} x' y' dV$$

$$I_{13} = \int_{V} \rho_{i} x' z' dV$$

$$I_{23} = \int_{V} \rho_{i} y' z' dV$$
(6.25)

where ρ is the density of the ice and V is its volume. For an ice mass which is symmetrical about the Gx'z' plane, I = I = 0; and for an ice mass which is symmetrical about the Gy'z' plane, I = I = 0.

The hydrostatic stiffness matrix components are given as:

where:

$$C_{33} = \rho g S_{w}$$
 $C_{34} = -\rho g S_{2}$
 $C_{35} = -\rho g S_{1}$
 $C_{44} = \rho g S_{22} + m g z_{B}'$
 $C_{55} = \rho g S_{11} + m g z_{B}'$

(6.27)

in which z' is the z' coordinate of the centre of buoyancy. Also S, S, S and S are the waterplane area and moments defined as:

$$S_{w} = \int dS$$

$$S_{1} = \int x' dS$$

$$S_{2} = \int y' dS$$

$$S_{11} = \int x'^{2} dS$$

$$S_{22} = \int y'^{2} dS$$

$$(6.28)$$

where the above integrals are taken over the waterplane. For an ice mass which is symmetrical about the x-z and y-z planes we have $S_1 = S_2 = 0$.

The viscous damping coefficients $\lambda(v)$ may be generally taken as zero except for heave, roll and pitch motions, corresponding to $\lambda(v)$, $\lambda(v)$ and $\lambda(v)$ and may be expressed in terms of empirical viscous damping ratios cj as:

$$\lambda_{jj}^{(v)} = 2c_{j}\sqrt{c_{jj}(m_{jj} + \mu_{jj})}$$
 (6.29)

(Note that the definition of c_j used here is based on values of μ at the oscillation frequency, whereas values based on the natural frequency are also used as an alternative.)

The solution to Eq. 6.23 can now be carried out to obtain the response amplitude operators which describe the ice mass responses for incident waves as required. With the values of ζ_k determined, the potential (Eq. 6.6) and thus the pressure field $(p=-\rho\partial\phi/\partial t)$, from the linearized Bernoulli equation) are explicitly known.

6.2.6 Wave-Drift Forces

In the drift motion calculation described in Section 6.3.2, wave drift forces, important in certain circumstances, are obtained from the diffraction solution described here. In the calculation of drift motions, the ice mass is assumed to have two degrees of freedom (surge and sway) and thus only the corresponding drift force components are required. These are obtained from formulae provided by Faltinsen and Michelsen (1974) which involve the source strengths f and the oscillatory motion amplitudes already determined.

6.2.7 Zero Frequency Added Masses

The drift motion calculation also depends on the zero frequency added masses of the ice with respect to surge and sway motions. These can be obtained from the diffraction calculation, now applied at a very low wave frequency. The Green's function need not be the same as that used in the wave diffraction calculation because a zero frequency limit is required. In this case, the simpler Green's function used in the current field computation may be employed (e.g. Garrison, 1978). However, the general procedure used to obtain the added masses exactly follows that outlined in this section.

6.2.8 Small Ice Masses

When the ice mass is relatively small (D/L less than about 1/10) it may be assumed not to affect the wave motion. This case is considerably simpler to treat since only the wave diffraction around the fixed structure need be considered in establishing the fluid flow. Furthermore, the computation need be carried out only once since it is now independent of the ice mass location.

The full theoretical development proceeds along the lines already discussed for a large ice mass, but various simplifications may be introduced. Only a single body, corresponding to the offshore structure is considered, so that S=S, and terms relating to S alone are absent. Secondly, since the structure is fixed, the forced potentials, corresponding to $k=1,\ldots,6$ are absent, and only the problem corresponding to k=7 need be treated. Only the 0xyz coordinate system is required to solve the diffraction problem and the location of the ice mass only enters the problem in determining the velocity components at its location. Once the potential ϕ_k has been obtained, the equations of motion are not solved directly as before, but rather the fluid velocity at the

location G of the ice mass may be obtained.

To account for resonance effects, the added mass and damping coefficients of the ice mass must be included into the equations of motion. The coefficients are determined from the same two body diffraction calculation as applied to larger ice masses.

6.3 Drift Motions

When the ice mass diameter to structure diameter ratio D /D is sufficiently small, the ice mass is considered to be in the nonuniform current field produced by the structure. The current field around the structure need then be computed and the trajectory of the ice mass then obtained by a time-stepping procedure applied to this non-uniform current field. If D /D is sufficiently large the structure has little influence on the current field around the ice mass. In this case, a uniform current field may be assumed and the presence of the structure on the current field ignored.

With respect to drift motions, the ice mass possesses three degrees of freedom (surge, sway and yaw). As an appropriate simplification, the drift in yaw is neglected. The extension to include yaw would require a computation of the flow field around the entire ice mass surface and applying this to a semi-empirical expression for the resulting torque on the ice mass. Due to flow separation effects, considerable uncertainty would be attached to this procedure.

Section 6.3.1 below describes the calculation of the current field around the structure which is required for relatively small values of D/D. If D/D is large, then the calculation can proceed directly with the equations of motion as described in Section 6.3.2.

6.3.1 Current Field for Small D /D i s

Since an input is only affected by the flow around the upstream side of the structure and since the ice mass is assumed to be relatively small, flow separation around both the structure and the ice mass are neglected. Potential flow theory is used, and the description given below largely follows that in Section 6.2 for the wave diffraction calculation.

A freely floating ice mass is situated in a uniform current which flows past a fixed large offshore structure, as depicted in As before, Oxyz forms a right-handed Cartesian 1. coordinate system with x measured horizontally and z measured upwards from the still water level. The fluid is assumed incompressible inviscid, and and the flow is assumed irrotational. The fluid motion is described by a velocity potential (defined such that the fluid velocity vector $\mathbf{u} = \nabla \phi$) which satisfies the Laplace equation within the fluid region. The potential \$\phi\$ is subject to boundary conditions on the free surface, the seabed, the structure submerged surface and the far field.

The velocity potential ϕ is expressed as:

$$\phi = \phi_{c} + \phi_{d} \tag{6.30}$$

where $\phi_{\mathbf{C}}$ is the potential associated with the incident current flow alone and $\phi_{\mathbf{C}}$ is the disturbance potential due to the presence of the structure. If the incident flow is represented as a uniform current of magnitude U flowing at an angle β to the x axis, then:

$$\phi_{C} = U(x\cos\beta + y\sin\beta) \tag{6.31}$$

The incident potential ϕ is thus known and ϕ is to be determined such that it also satisfies the Laplace equation and that the boundary conditions are satisfied.

The body surface boundary condition requires that there is no flow velocity normal to the submerged structure surface S and may be written as:

$$\frac{\partial \phi}{\partial n} (\underline{\mathbf{x}}) = \frac{1}{2} b(\underline{\mathbf{x}}) \qquad \underline{\mathbf{x}} \text{ on } S$$

$$b_{\mathbf{k}} (\underline{\mathbf{x}}) = -2 \frac{\partial \phi}{\partial n} c(\underline{\mathbf{x}}) \qquad (6.32)$$

and n denotes distance in the direction of the unit normal vector n directed normally out from S.

The disturbance potential ϕ may be represented as a distribution of point sources around the submerged structure surface S. Thus the potential $\phi_d(\underline{x})$ at a general point \underline{x} within the fluid may be expressed as:

$$\phi_{\rm d}(\underline{\mathbf{x}}) = \frac{1}{4\pi} \int_{\rm S} \mathbf{f}(\underline{\xi}) \mathbf{G}(\underline{\mathbf{x}}, \underline{\xi}) d\mathbf{S}$$
 (6.33)

where $f(\underline{\xi})$ is a source strength distribution function, $\underline{\xi}$ is a point on the surface S over which the integration is performed, and $G(\underline{x},\underline{\xi})$ is a Green's function for the general point \underline{x} due to a source of unit strength located at $\underline{\xi}$.

G is itself chosen to satisfy the Laplace equation and the various boundary conditions. The current is assumed to be sufficiently slow so that free surface effects can be neglected

and the free surface condition then corresponds simply to the requirement that the vertical velocity there is zero. The Green's function thus corresponds to that of a point source with reflection terms to account for the seabed and free surface conditions, and is known (e.g. Garrison, 1978).

The source strength function f is required and may be determined by applying the body surface boundary condition. This condition, Eq. 6.32, yields a surface integral equation for f:

$$-f(\underline{\mathbf{x}}) + \frac{1}{2^{\pi}} \int_{S} f(\underline{\xi}) \frac{\partial G}{\partial n} (\underline{\mathbf{x}}, \underline{\xi}) dS - b(\underline{\mathbf{x}})$$
 (6.34)

where \underline{x} now lies on S at the point where the boundary condition is applied, n is measured from \underline{x} and the integration is carried out over the points $\underline{\xi}$.

In a numerical solution to Eq. 6.34 the surface is discretized into N small facets with the function f taken as uniform over each facet. The surface integral equation is then approximated as a set of linear algebraic equations for F = f(x):

$$\sum_{j=1}^{N} A_{ij} f_{j} = b_{i} \qquad i=1,...,N$$
(6.35)

where $b = b(\underline{x})$ and \underline{x} is the value of \underline{x} at the centre of the ith facet. The matrix coefficients may be approximated as:

$$A_{ij} = \begin{bmatrix} \frac{\Delta S}{2\pi} \frac{\partial G}{\partial n} & (\underline{x}_{i}, \underline{\xi}) & i \neq j \\ -1 & i = j \end{bmatrix}$$
 (6.36)

where ΔS is the area of the j-th facet. With the source strengths f evaluated, the velocity potential is obtained by a

discretized form of Eq. 6.33. The velocity components obtained by differentiating Eq. 6.33 with respect to the corresponding directions and then applying a similar discretized procedure. The current velocity components u and u in the x and y directions respectively are given as:

$$u_{i} = \frac{\partial \phi}{\partial n_{i}} (\underline{\mathbf{x}}) = \frac{\partial \phi}{\partial n_{i}} c(\underline{\mathbf{x}}) + \sum_{j=1}^{N} b_{j} f_{j} \qquad i=1,2 \qquad (6.37)$$

where

$$\mathbf{B}_{\mathbf{i}\mathbf{j}} = \frac{\Delta S}{4\pi} \mathbf{j} \frac{\partial G}{\partial \mathbf{n}_{\mathbf{i}}} (\underline{\mathbf{x}}, \underline{\xi}_{\mathbf{j}}) \tag{6.38}$$

and n_{i} is here taken as the x and y directions in turn.

6.3.2 Equations of Motion

The forces acting on the ice mass are comprised of the current drag, the wave drift force, the wind drag and the combined force associated with the Coriolis force and the effect of sea surface slope. For short term motion predictions in the vicinity of a structure as required here, the current drag and possibly the wave drift force are expected to be the most significant and are included while the remainder are neglected.

The horizontal drag force vector $\underline{\mathbf{F}}^{(d)}$ involves the local current velocity $\underline{\mathbf{u}}$ and the ice mass velocity $\underline{\mathbf{u}}$ and may be expressed as:

$$\frac{\mathbf{f}}{\mathbf{f}} = \frac{1}{2}\rho \mathbf{A} \mathbf{C}_{\mathbf{d}} \mathbf{u} | \mathbf{u} |$$
 (6.39)

where \underline{u} is the current velocity vector relative to the ice mass and is given as $\underline{u} = \underline{u} - \underline{u}$, C is a drag coefficient and A is the projected area of the ice mass on to the plane perpendicular to \underline{u} . Equation 6.39 may be expressed in terms of the drag force components in the x and y directions as:

$$F_{i} = \frac{1}{2} \rho AC_{d} u_{i} | \underline{u} | i=1,2$$
 (6.40)

where the subscripts 1 and 2 denote the components in the x and y directions respectively. The wave drift force \underline{F} may be computed from the diffraction calculation as already described in Section 6.2.6.

The equations for surge and sway may now be expressed as:

$$(m + \mu_i)\dot{u}_i = F_i + F_i$$
 i=1,2 (6.41)

where μ_1 , μ_2 are the zero frequency added masses in the x and y directions respectively and may be obtained as indicated in Section 6.2.7. The equations of motion are conveniently written as:

$$\dot{u}_{i} = a_{i}$$
 $i=1,2$ (6.42)

where

$$a_{i} = (F_{i}^{(d)} + F_{i}^{(w)})/(m + \mu_{i})$$
 (6.43)

Equation 6.42 represents a pair of nonlinear equations for u and may be solved by a suitable time-stepping procedure. Many such methods are available, including the Runge-Kutta methods and others. As a relatively simple but reliable method the Adams-

Bashforth two-step formula may be used and gives the solution at an advanced time $t+\Delta t$ in terms of the solution at previous time steps t and $t-\Delta t$ as:

$$u_{i}(t+\Delta t) = u_{i}(t) + \frac{1}{2}\Delta t[3a_{i}(t) - a_{i}(t-\Delta t)]$$
 (6.44)

In applying Eq. 6.44, the right-hand sides contain quantities at times t and t - Δt which are known from previous iterations so that the ice mass velocity u at the new time t + Δt may be obtained.

Once the drift velocity has been found, the new location may also be determined by a time-stepping procedure applied in a similar way to obtain the new location of the ice mass, described by the coordinates x, y. Thus:

$$x_{G}(t+\Delta t) = x_{G}(t) + \frac{1}{2}\Delta t[3u_{1}(t) - u_{1}(t-\Delta t)]$$

$$y_{G}(t+\Delta t) = y_{G}(t) + \frac{1}{2}\Delta t[3u_{2}(t) - u_{2}(t-\Delta t)]$$
(6.45)

6.4 Computer Procedure

6.4.1 Computational Considerations

The accuracy of the diffraction calculation is strongly influenced by the discretization of the submerged ice mass and structure surfaces into a suitable number of facets. The usual representation of the surfaces is by plane polygonal facets. To ensure that the surfaces are divided up sufficiently finely, facet diameters should be less than about 1/8 of the incident wave length and also sufficiently small so as to represent the surface configuration realistically. One or two hundred facets are typically required.

diffraction calculation is computationally rather expensive, with computer storage and cost considerations limiting the number facets used. For a single diffraction calculation computation involves the calculation of a corresponding set coefficients, which include lengthy expressions based on the Green's function. and a solution of full the equation. Since the interaction between the ice mass and structure is of paramount interest, the diffraction calculation to be repeated for different ice mass locations along trajectory, but it need not be repeated at every time step of the drift motion calculation. Rather it is repeated every several time steps and intermediate values of zero frequency added mass and wave drift force, required for the drift calculation, obtained by a simple extrapolation procedure.

Numerical instabilities in the diffraction procedure generally expected since considerable experience has been in the context of motion response calculations for floating vessels, including interference effects between neighbouring Some numerical difficulties may arise when the facets on the ice mass and structure surfaces are less than about one facet diameter apart (Section 7). These can be avoided for particular situations by a judicious choice of facet size and the time-stepping procedure applied to the equations of motion is of a general form which has been rather extensively in the past, and associated numerical instabilities are not expected.

6.4.2 Program Computations

The proposed computer procedure is indicated in the flow chart given in Figure 2. The facet data required for the diffraction and wave field solutions are prepared by a pre-processing program. The main procedure involves an initial diffraction

calculation to obtain initial values of the wave-induced motions, the wave drift forces and the zero frequency added masses. If the structure diameter to ice mass diameter ratio is sufficiently large the current field around the structure is then computed and the initial current velocity at the ice mass location is determined.

The program then enters a time loop with calculations performed at successive time steps to trace the trajectory of the ice mass. At each time step the equations of motion of the ice mass are used to obtain the updated drift velocity of the ice mass and its new location. The current velocity at this new location is then calculated if necessary (i.e. if D/D is sufficiently large). Every few time steps, the diffraction solution is also carried out to obtain the wave-induced motions, the wave drift force and the zero frequency added masses corresponding to this new location. The time loop is then repeated to provide an updated drift velocity and location of the ice mass.

A fairly small time step size is needed in order to reproduce a sufficiently accurate drift trajectory of the ice mass. The variation of oscillatory motions, wave drift forces and hydrodynamic coefficients along the trajectory is not expected to be required to the same level of precision. The major computational effort corresponds to the heavily bordered blocks in Figure 2. Of these, the diffraction calculation at successive time steps is the most lengthy, and may be calculated with a coarser time interval than that used for the drift equations of motion calculation. This refinement is indicated in the figure.

7.0 PARAMETRIC AND SENSITIVITY TESTS

Based on the theoretical and numerical considerations outlined in Section 6, a computer model was developed, and tests were undertaken to demonstrate the capabilities of the model. This section presents the results of the parametric and sensitivity tests.

7.1 Test Conditions

A series of numerical tests are described which assess the influences of variations in ice mass size, wave direction, initial ice mass location, current magnitude and water depth, and examines the influence of wave drift forces. A considerable number of parameters are required as input to the computer model and, due to time and budget restraints, many are held constant. The structure is represented as a fixed, surface-piercing, vertical, circular cylinder of diameter 100 m extending to the seabed. The three ice masses are represented as floating circular cylinders of diameters 100 m, 50 m and 10 m, all with drafts equal to the corresponding radii. The dimensions and dynamic properties of the three ice masses are summarized in Table 1.

Environmental parameters for the tests are:

Wave height: 10 m
Wave period: 15 s
Current direction: 0
Drag coefficient: 1.0
Viscous damping ratio: 0.05

Current magnitude: 1.0 m/s (except test 9, where

U = 0.4 m/s)

Water depth: 100 m (except test 10, where d = 60 m)

TABLE 1. ICE MASS PROPERTIES

	Ice Mass		
•	A	В	С
D _i (m)	100	10	50
h (m)	50	5	25
(m)	5.9	0.6	2.9
\mathfrak{m}/ρ (\mathfrak{m}^3) 39	92700	392.7	49087
$I_1/m = I_2/m (m^2)$	385.3	8.85	221.3
$I_3/m (m^2)$	250	12.5	312.5
z _G (m) - 23	2.06	- 2.21	- 11.03
z _B (m) - :	25	- 2.5	- 12.5
s (m ²) 78	854	78.5	1963.5
$s_{11} = s_{22} (m^4)$ 4	.91 X 10 ⁶	491	307 X 10 ³
T _n (sec)	18	6	12

TABLE 2. TEST CONDITIONS

TEST	ICE MASS	WAVE DIR a(°)	INITIAL OFFSET Y _G (m)	WAVE DRIFT FORCE	OTHER/COMMENT
1	A	0	0	Incl	Symmetrical
2	A	0	25	Incl	Initial offset
3	A	5	0	Incl	Oblique wave dir.
4	A	5	25	Incl	Initial offset and oblique wave dir.
5	В	0	0	Incl	Common hasi wa 1
5A	В	0	0	Excl	Symmetrical
6	В	0	25	Incl	T. 14.1.3
6 A	В	0	25	Excl	Initial offset
7	В	0	1	Incl	
7 A	В	0	1	Excl	Very small initial offset
8	В	5	0	Incl	
8A	В	5	0	Excl	Oblique wave dir.
9	A	5	25	Incl	Current = 0.4 m/s
10	A	0	25	Incl	Depth = 60 m
11	С	0	25	Incl	<pre>Initial offset; Ice mass C</pre>
. 12	A	Coorde			
13	В	specifi	ied locatio	ons close	to structure

lists the conditions for the seventeen test runs (numbered 1 through 13, and 5A through 8A. Tests 1-4 involve the large ice mass A with different combinations of initial offset y and wave direction $\boldsymbol{\alpha}$. A similar set of tests were undertaken for ice mass B in tests 5-8. In the case of the small ice mass, however, the wave drift force has been found to be a critical factor in the drift motion and thus pairs of tests have been carried out so as to include (tests 5, 6, 7, 8) and exclude 8A) the wave drift force. 6A, 7A, Tests 7 and 7A include a small initial offset $y_c = 1 \text{ m}$ to test the possibility of an instability in the drift motion. Test 9 corresponds to the same conditions as test 4 except that the current magnitude reduced, U = 0.4 m/s rather than 1.0 m/s. Test 10 corresponds to the same conditions as test 2 except that the water depth reduced, d = 60m rather than 100 m. Test 11 treats intermediate size ice mass C with initial offset y = 25wave direction $\alpha = 0$. Finally, tests 12 and 13 have been carried out with the large and small ice masses relatively close structure to examine the possible variation of frequency added mass as the gap between the ice mass and structure is reduced.

The numbers of facets used to discretize the ice masses and structure are summarized in Table 3. In the table, N and N are respectively the number of circumferential and depth divisions used to carry out the discretization. The values given correspond to a maximum facet side length of about 25 m for most tests, except that this length is smaller in tests 12 and 13. The numerical results are unreliable when the gap between the ice mass and structure or seabed is less than the facet size.

The initial drift velocity of the ice mass is generally chosen to give it a zero or small initial x-ward drift acceleration so that the drag force counteracts the wave drift force. The ice mass undergoes a drift acceleration if it is initially drifting at the current velocity because the current drag is then zero and the

TABLE 3. DISCRETIZATION OF STRUCTURE AND ICE MASSES

TEST	ICE	S	TRUCTURE		Ī		
	MASS	N	N _c	N _d	N	N _C	N _d
1-4, 9	Ą	48	12	4	37	12	2
5-8, 5A-8A	В	48	12	4	25	8	2
10	A	48	12	4	65	16	3
11	C.	48	12	4	37	12	2
12	A	108	18	6	73	18	3
13	В	120	20	6	25	8	2

TABLE 4. PERCENTAGE INCREASE IN ZERO FREQUENCY

ADDED MASS WITH GAP SIZE FOR ICE MASS A

		GAP SIZE	
	15 m	10 m	5 m
Surge	4.3	9.4	_
Sway	1.8	3.3	-3.1

wave drift force is unbalanced. The initial x-ward locations of the ice masses were taken to correspond to x = -300 m for ice mass A and to x = -200 m for ice masses B and C.

7.2 Test Results

The computer model output provides (1) the specified input parameters; (2) parameters initially calculated, including various wave frequency parameters, and mass and stiffness matrices of the ice mass; and (3) at each time step, the ice mass velocity and location and the current velocity at the ice mass location. For those time steps at which the diffraction computations are carried out, the output also provides the exciting forces (magnitude and phase), added masses, damping coefficients, response amplitude operators (magnitude and phase), wave drift forces, and the zero frequency added mass (a sample output is shown in the Appendix).

Since a large number of parameters are generated for each test, only the major ones are presented to illustrate the fundamental features of the different motions predicted. The corresponding results of tests 1-11 are shown in Figures 3 - 18. The figures show the ice mass trajectory, the x-ward and y-ward drift velocity components u and u, the zero frequency added masses μ_1 and μ_2 , the radiation damping coefficients λ_{11} , λ_{2} and λ_{33} , and the motion amplitudes ζ_1 , ζ_2 , and ζ_3 , all as functions of the x-ward location of the ice mass, x $_{G}$. (Note that y and x are plotted to different scales so that the trajectory shown is generally distorted.) In the figures, the following dimensionless forms of the hydrodynamic coefficients are used:

$$\mu_{\mathbf{i}}' = \frac{\mu_{\mathbf{i}}}{\rho D_{\mathbf{i}}} 3 \tag{7.1}$$

$$\mu'_{ii} = \frac{\mu_{ii}}{\rho D_{i}} 3$$

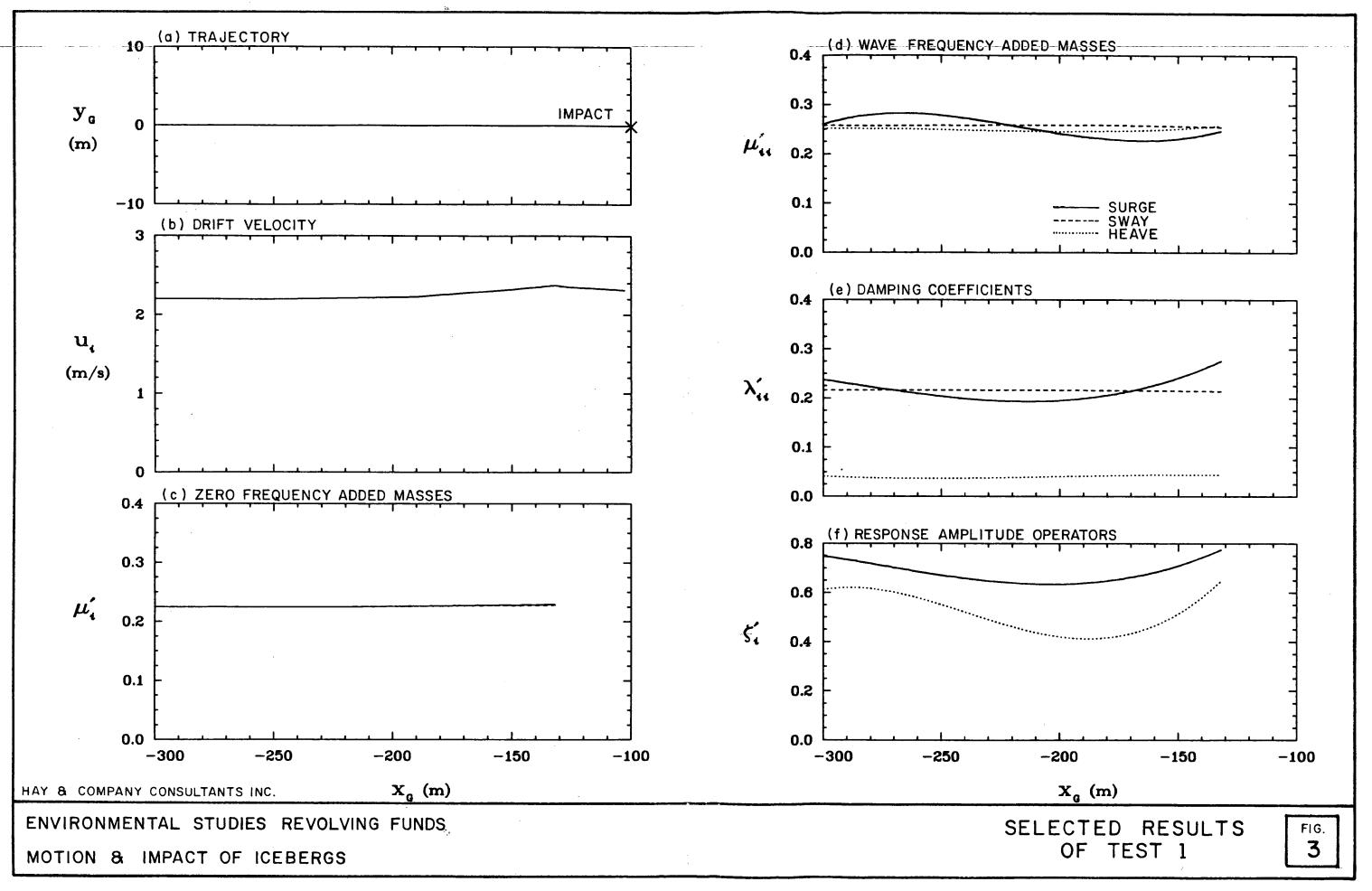
$$\lambda_{ii}' = \frac{\lambda_{ii}}{\rho \omega D_{i}} 3$$
 (7.1) cont'd.

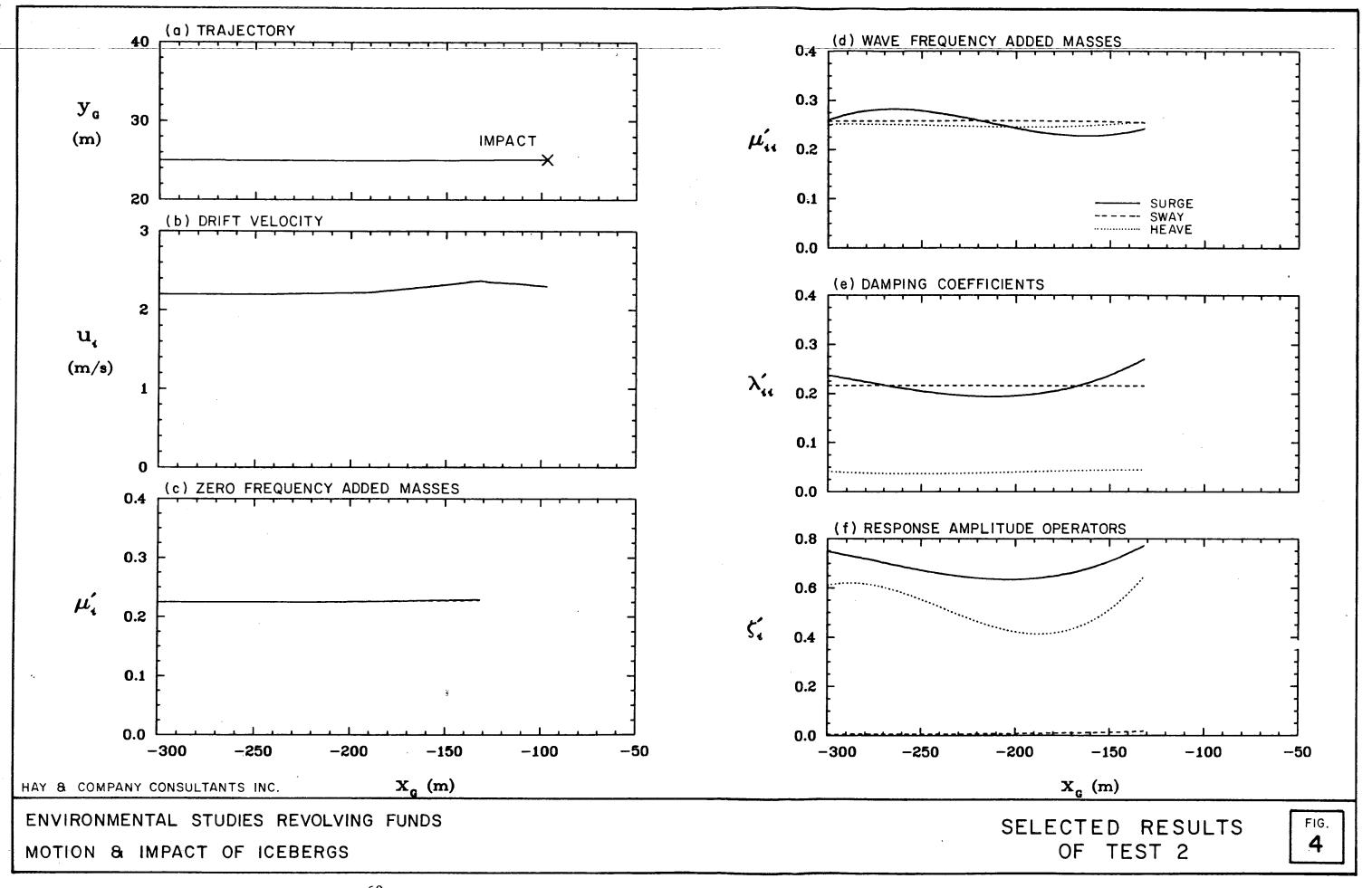
$$\zeta_{\mathbf{i}}' = \frac{\zeta_{\mathbf{i}}}{H/2}$$

Here ρ is the water density, D is the ice mass diameter, H is the wave height, and i = 1, 2, $\frac{1}{3}$ denote the surge, sway and heave components respectively. ζ corresponds to the response amplitude operators (RAO's) of the ice mass and is defined with respect to the still water level.

The zero frequency added mass in heave is theoretically infinite for the finite depth cases treated here and so these are absent from the figures.

The results of test 1 are shown in Figure 3. Since the test conditions are symmetrical, the ice mass does not deviate from its initial direction so that y remains at zero (Figure 3(a)). Figure 3(b) shows the x-ward drift velocity u_1 to vary only slightly because the ice mass is relatively large in size, and the y-ward component u_2 remains at zero. (The abrupt changes in drift velocity shown here and in subsequent results associated with a relatively coarse time interval used in the diffraction calculations.) Figure 3(c) indicates that the zero frequency added masses increase only when the gap between the ice mass and structure is of order 0.1 of the ice mass diameter. Thus, only immediately prior to impact is the effect of proximity to the structure important relative to the zero frequency added stated in Section 6.4.2, at such separations, masses. numerical limitations arise from the facet size. This effect is considered subsequently in tests 12 and 13. Finally, Figures show the wave frequency added masses, coefficients and response amplitude operators all as functions of





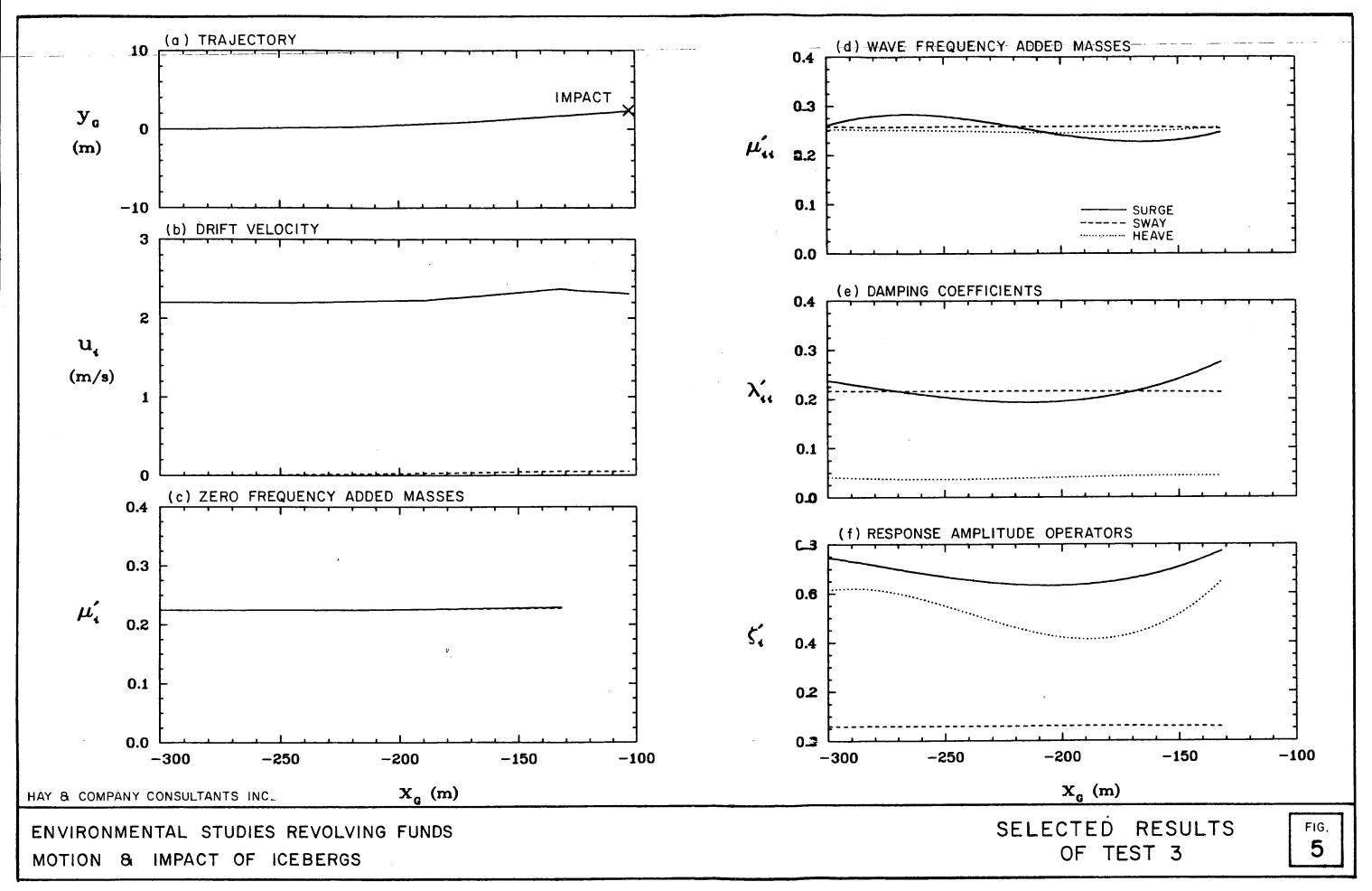
 \mathbf{x} . These all show cyclic variations with \mathbf{x} , with a length scale corresponding to the incident wave length. This cyclic behaviour is associated with the interference effect of the structure on the wave field experienced by the ice mass. As with all the tests of zero initial offset and wave direction, the sway RAO is zero.

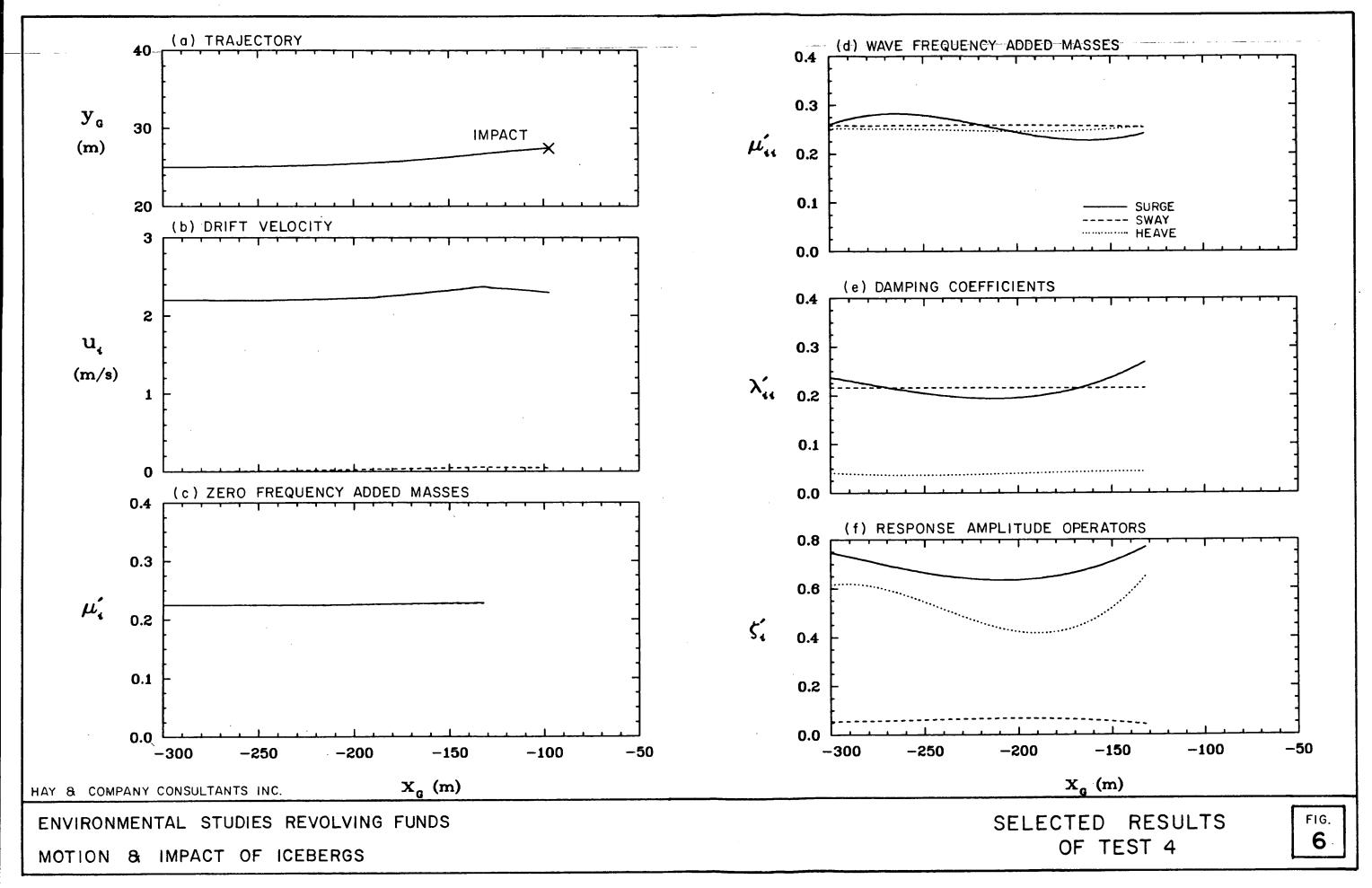
Test 2 (Figure 4) corresponds to an initial offset $y=25\,\mathrm{m}$ and a wave direction $\alpha=0$. The results are very similar to those of test 1. Figure 4(a) indicates that because of the relatively large ice mass size, the ice mass continues drifting in its initial direction so that y remains at 25 m. The variation of hydrodynamic coefficients with x is virtually the same as in test 1, except that there is now a small sway component to the RAO caused by the asymmetrical interference effect of the structure.

Test 3 (Figure 5) corresponds to a zero initial offset and an oblique wave direction $\alpha=5^\circ$. Because of the influence of wave drift forces, the trajectory is now slightly curved with a y_G value of about 2 m at impact, and a small y-ward component to the drift velocity developing. Again, results for the hydrodynamic coefficients are similar to those of test 1, but now the sway response amplitude operator (RAO) is more noticeable because of the oblique wave direction.

Test 4 (Figure 6) corresponds to an initial offset y = 25 m and an oblique wave direction $\alpha = 5$. There is a similar trend to the trajectory and drift velocity as in test 3, except that this is now with respect to the initial value of y. As expected, the hydrodynamic coefficients are very similar to those in the previous tests.

Test 5 (Figure 7) treats a small ice mass with symmetrical initial offset and wave conditions, y = 0 m and $\alpha = 0$. The trajectory remains along y = 0 m as required by symmetry.





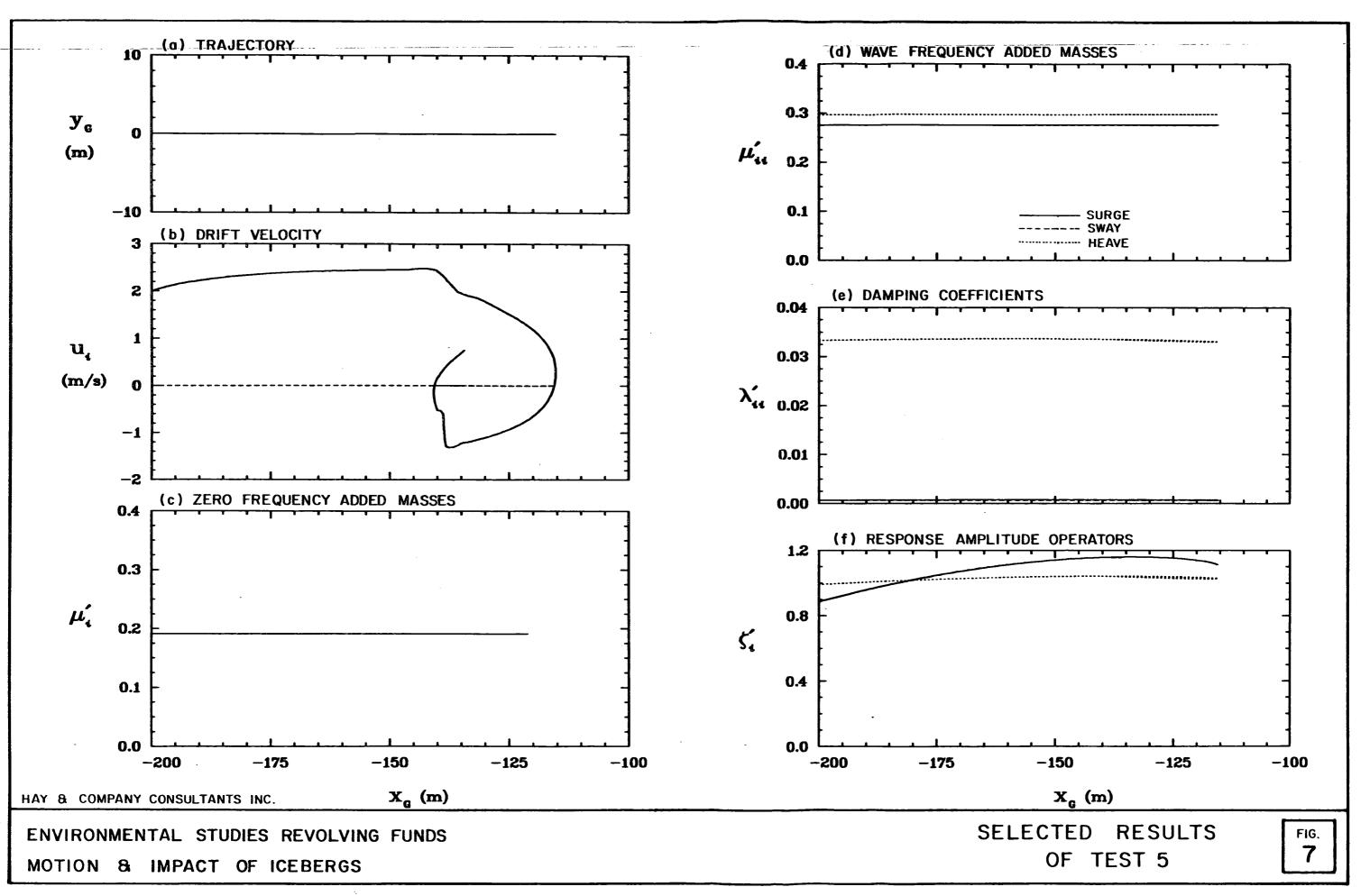


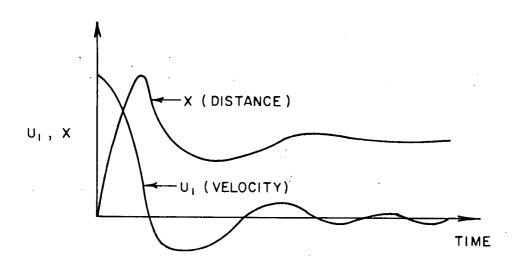
Figure 7(b) shows that the drift velocity first increases, then decreases and then increases again, such that ice mass reaches a maximum value of x, then reverses direction travelling away from the structure, and eventually undergoes decaying oscillations in its x-ward location.

This somewhat surprising result may be explained by examining the current drag and the wave drift force. The current drag opposes the relative motion between the ice mass and the current. For this ice mass the wave drift force varies cyclically with distance from the structure, such that it is negative over certain ranges of x and positive otherwise. The negative drift force (i.e. acting in an opposite direction to that of the wave propagation) is associated with the interference effect from the structure. Previous studies (e.g. Loken, 1981, van Oortmerssen, 1981) have indicated that cyclic variations in wave drift force with distance from a body (or with wave frequency) do occur and negative wave drift forces have been encountered. The effect is more pronounced here because of the relatively small ice mass.

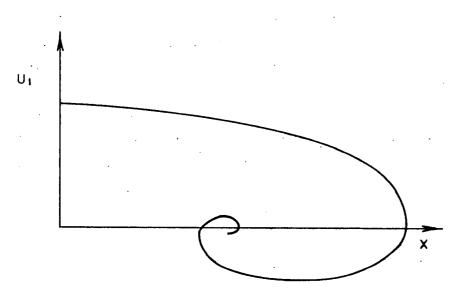
The effect of this kind of drift force behaviour on the drift motion may be examined through a simple drift model in which the wave drift force varies cyclically with distance from the structure. The drift equation of motion may be simplified to the form:

$$\dot{u}_1 = A \cos(kx) + B (u_1 - U) |u_1 - U|$$

where A and B are parameters describing the relative magnitude of the wave drift and drag forces, k is a wave number and U is the (uniform) current velocity. The solution to this equation exhibits cyclic variations in u and gives rise to negative values of u provided that the ratio A/B is sufficiently large. A sketch of a typical solution is shown in Figure 8. Figure 8(a) shows u and x as functions of time, and Figure 8(b) shows u as a function of x. These results are of the same general form to



(a) DRIFT VELOCITY AND LOCATION VARIATION WITH TIME.



(b) DRIFT VELOCITY VARIATION WITH LOCATION.

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MOTION & IMPACT OF ICEBERGS

RESULTS OF DRIFT MODEL

FIG.

8

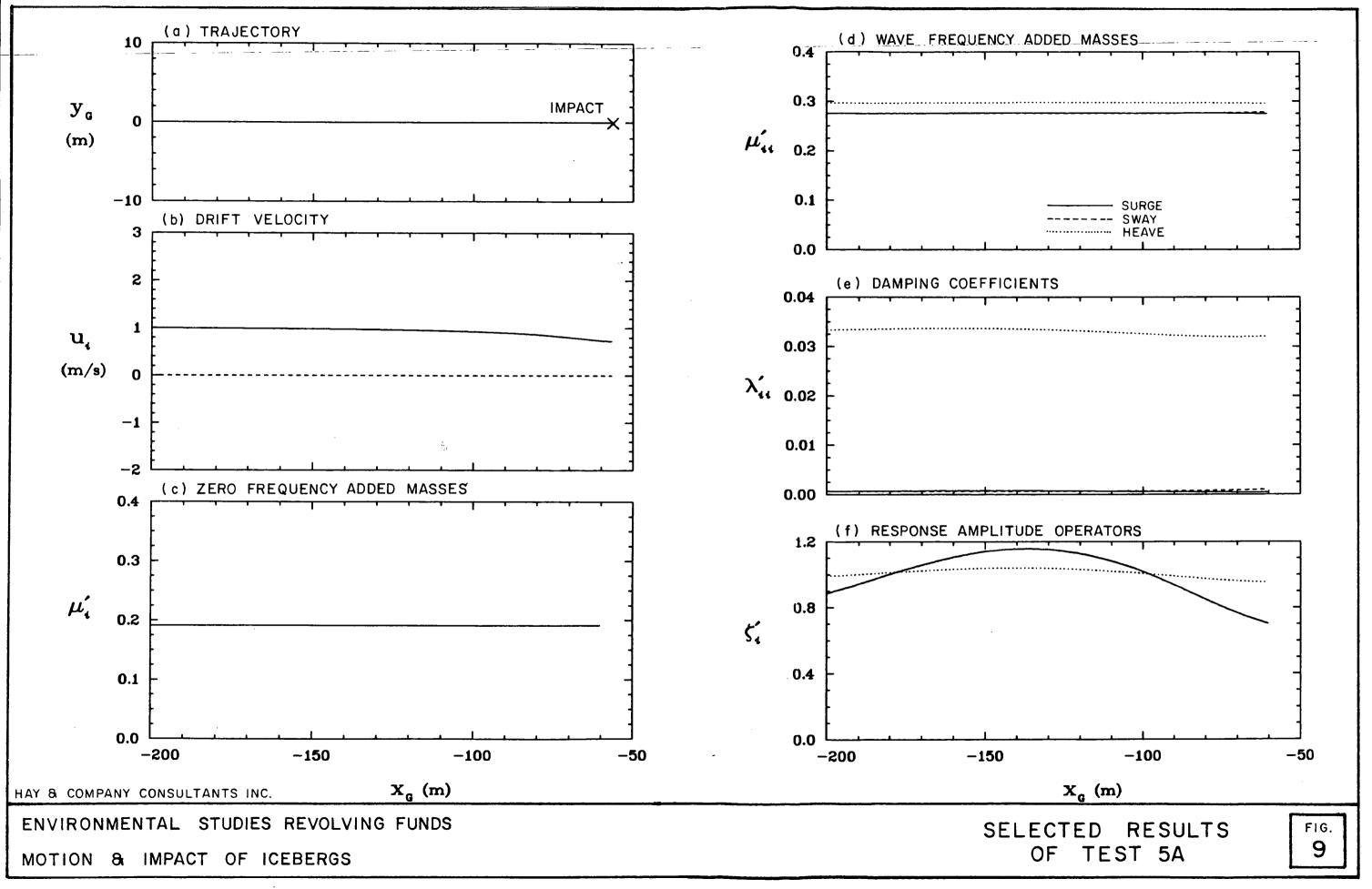
those in Figure 7b.

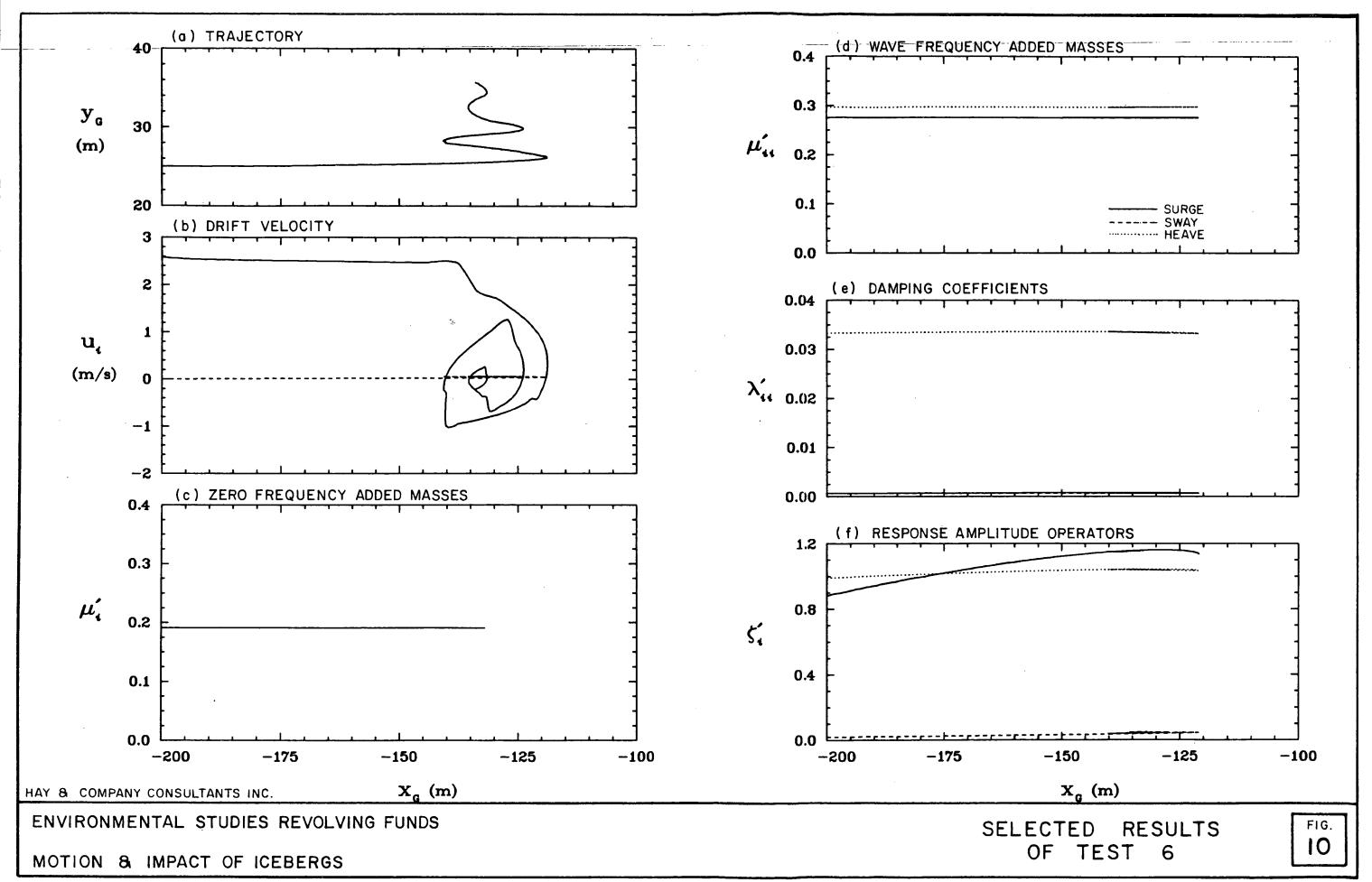
The hydrodynamic coefficients, Figures 7(c)-(e) do not show a noticeable variation with distance from the structure. The range of x shown is restricted because of the drift reversal. The surge exciting force, which is not shown, varies more strongly with x and thus the surge RAO shows a more noticeable variation with x. The sway RAO is zero because of symmetry.

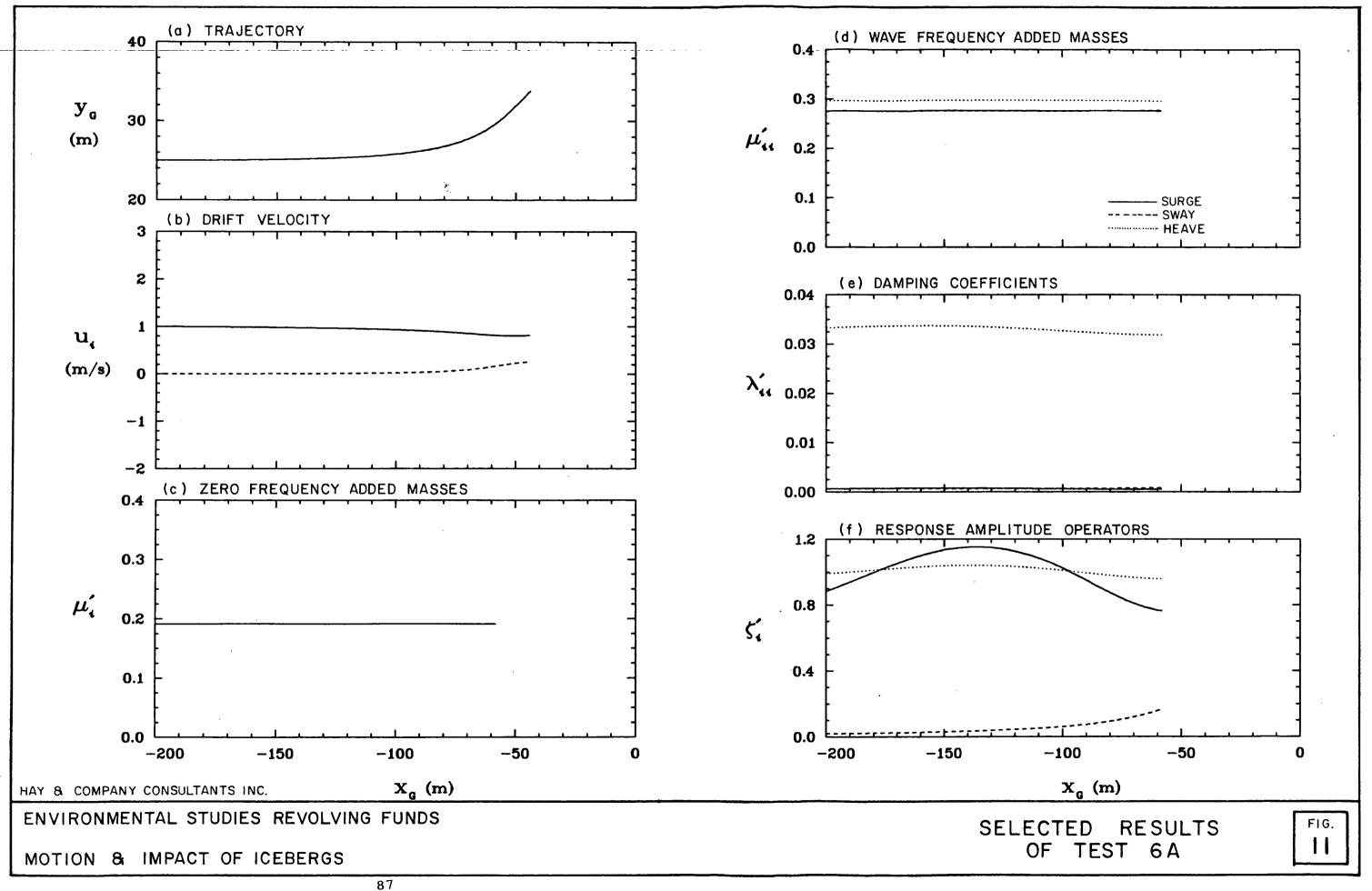
To investigate further the influence of the wave drift force, test 5A (Figure 9) corresponds to the same conditions as test 5, except that the wave drift force is now excluded from the drift equations of motion and the initial drift velocity is modified to provide for no initial x-ward acceleration. In this case the reversal of the drift motion is no longer present and the ice mass continues along $y_G = 0$ with a slightly decreased drift velocity until impact occurs. The hydrodynamic coefficients do not change strongly with x_G , but as the ice mass approaches the structure more closely than in test 5, the cyclic variation in the surge RAO is now apparent.

Test 6 (Figure 10) corresponds to an initial offset $y_G = 25 \text{ m}$, and shows the same behaviour in the x-ward component of drift velocity as in test 5. The continuous increase in y_G is due to the non-uniform current flow around the structure. This combines with the reversing x-ward drift to produce the trajectory shown in Figure 9(a). As in test 5, the hydrodynamic coefficients do not change strongly with location, except that the surge RAO does increase as the ice mass approaches the structure. A small sway RAO muses from the sway exciting force associated with the asymmetric interference effect of the structure.

Test 6A (Figure 11) shows the results for the same conditions as test 6, except that the wave drift force is omitted, and the initial drift velocity is correspondingly altered. As in test 5A, the trajectory and drift velocity now indicate no motion





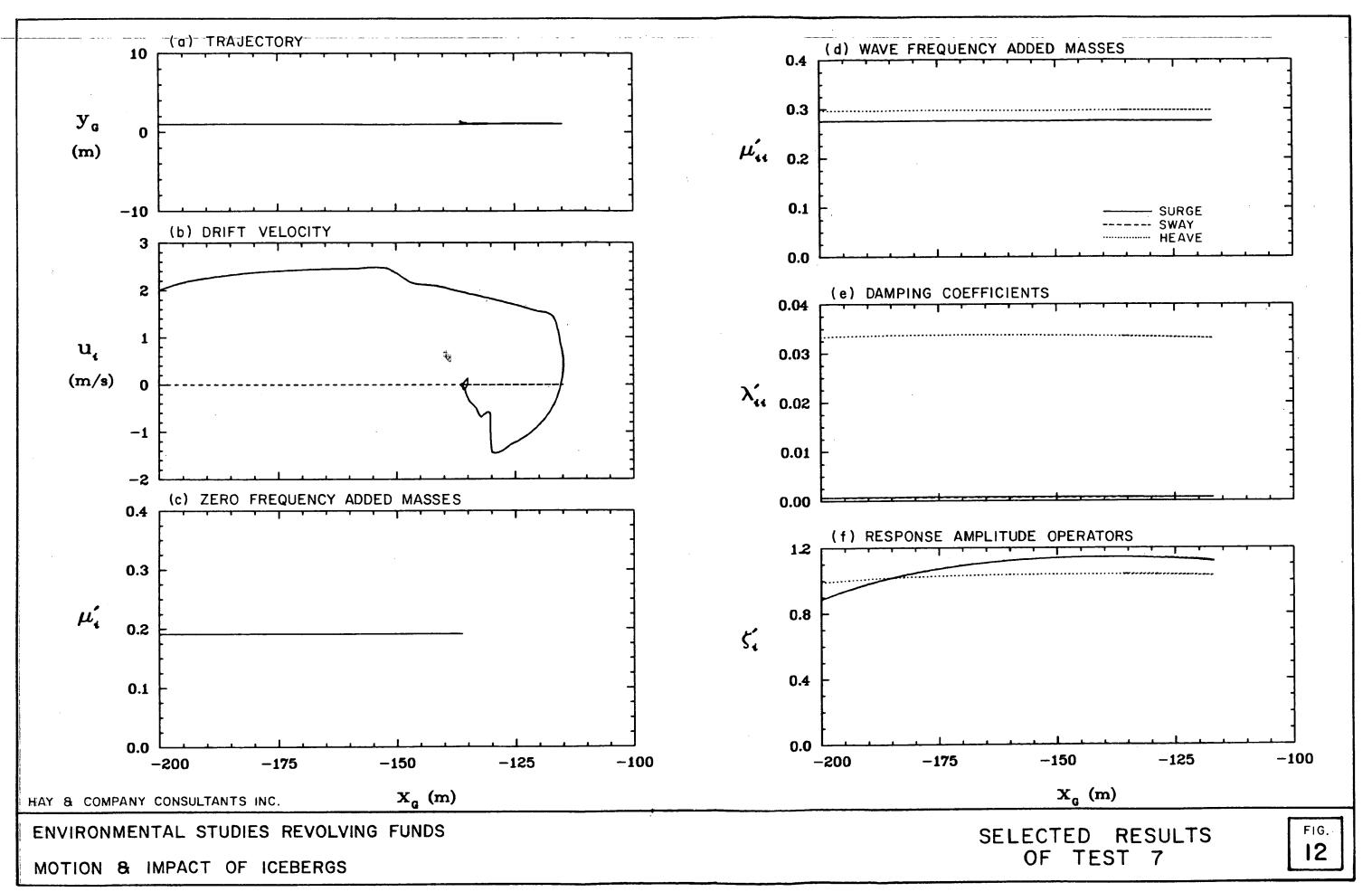


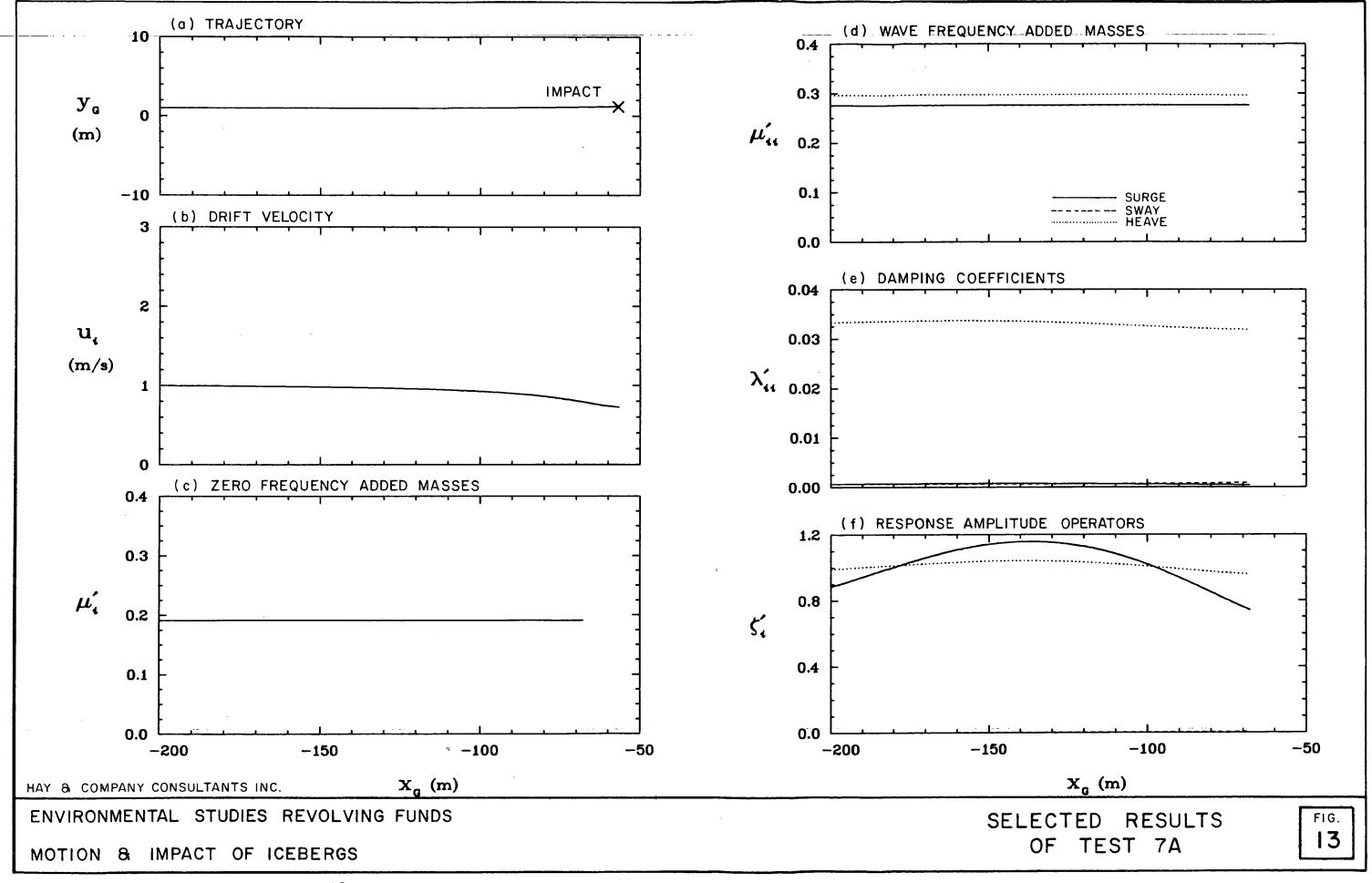
reversal. Similarly, as the ice mass approaches the structure more closely than in test 6, the variation in surge RAO with x is again apparent, and the sway RAO increases as the gap reduces.

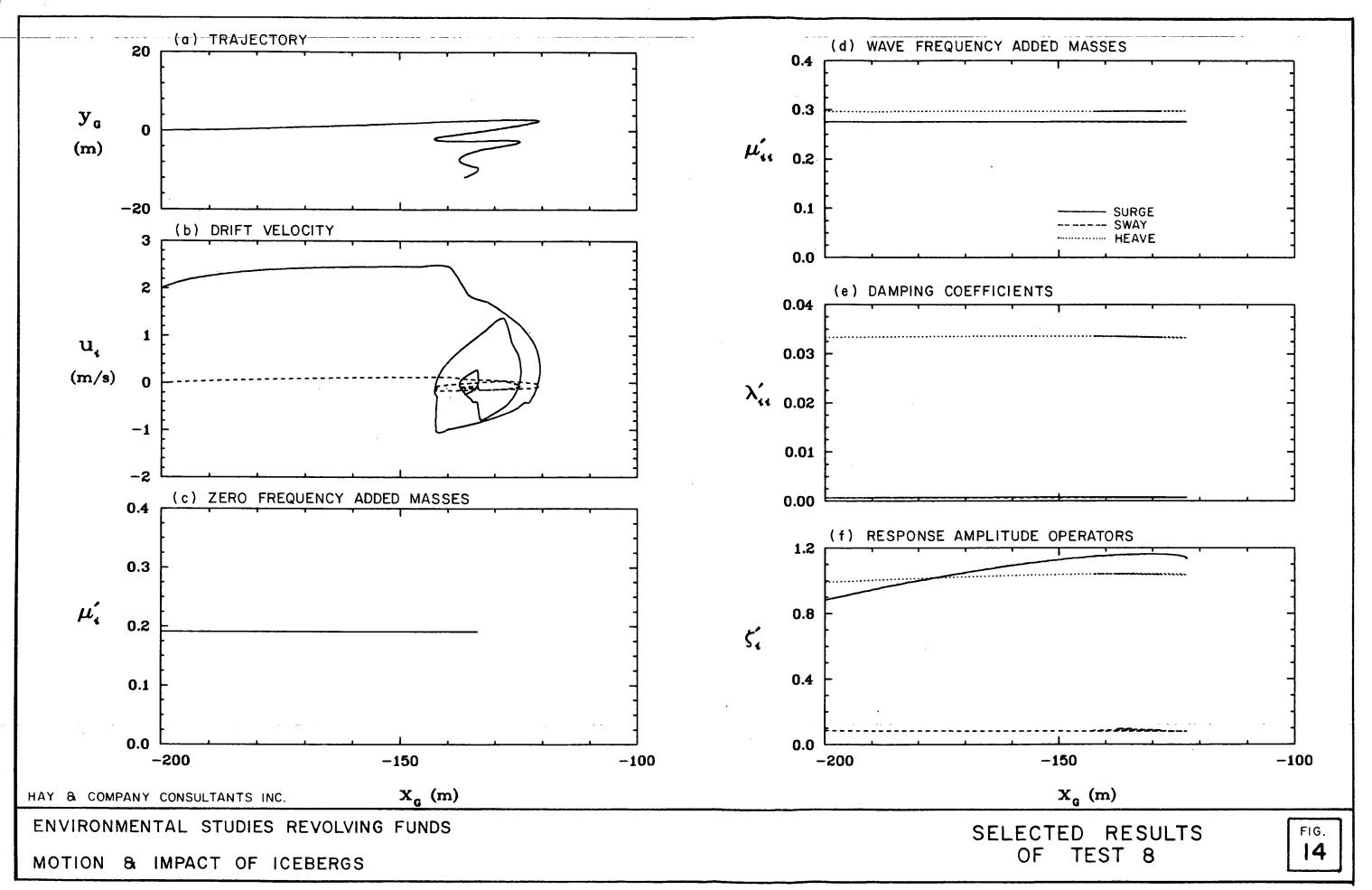
Test 7 (Figure 12) corresponds to an initial offset $y_G = 1 \text{ m}$ instead of the 25 m value used in test 6. y_G does not increase significantly as the ice mass approaches the structure because the y-ward component of the current velocity is very small. Test 7A (Figure 13), with the wave drift force excluded, shows again that y_G hardly changes and that the x-ward component of drift velocity u_G decreases slightly as the ice mass approaches the structure. The hydrodynamic coefficients behave in almost the same way as in test 6A, with a very small sway component now developing because of the slight asymmetry in the drift trajectory.

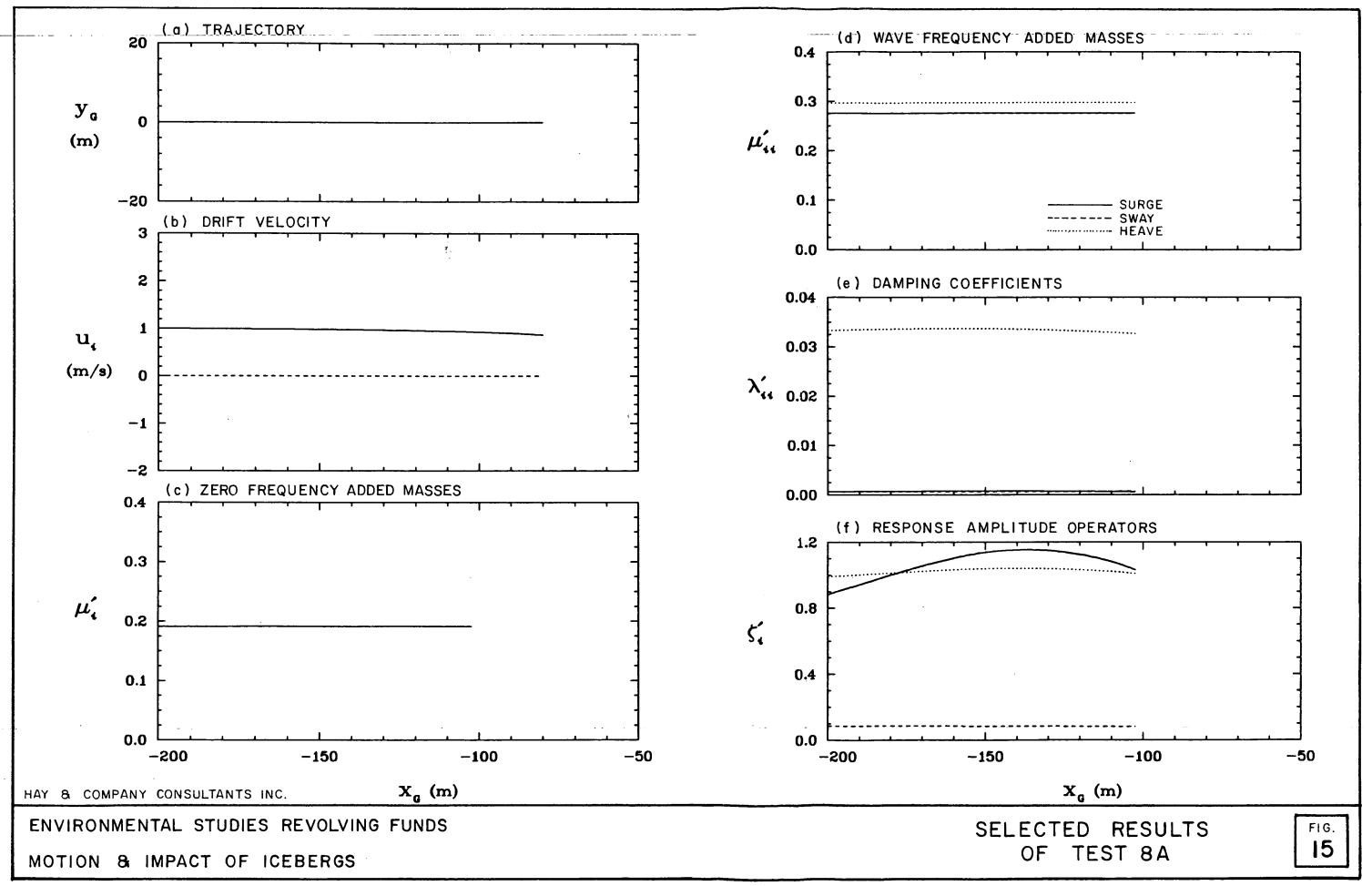
Test 8 (Figure 14) corresponds to a zero initial offset and an oblique wave direction $\alpha = 5$ °. Since the wave drift force acts primarily in the wave direction, y starts to increase from its initial zero value; when the wave drift force later becomes negative for certain ranges in x, y then decreases and hence the trajectory develops as shown in Figure 14(a). With the wave drift forces excluded from the drift equations of motion (test 8A, Figure 15), no y-ward excitation to the drift motion arises and hence y remains at zero. The x-ward drift velocity increases as in test 6. The surge RAO shows some cyclic variation with x as before, and there is a small sway RAO due to the wave direction.

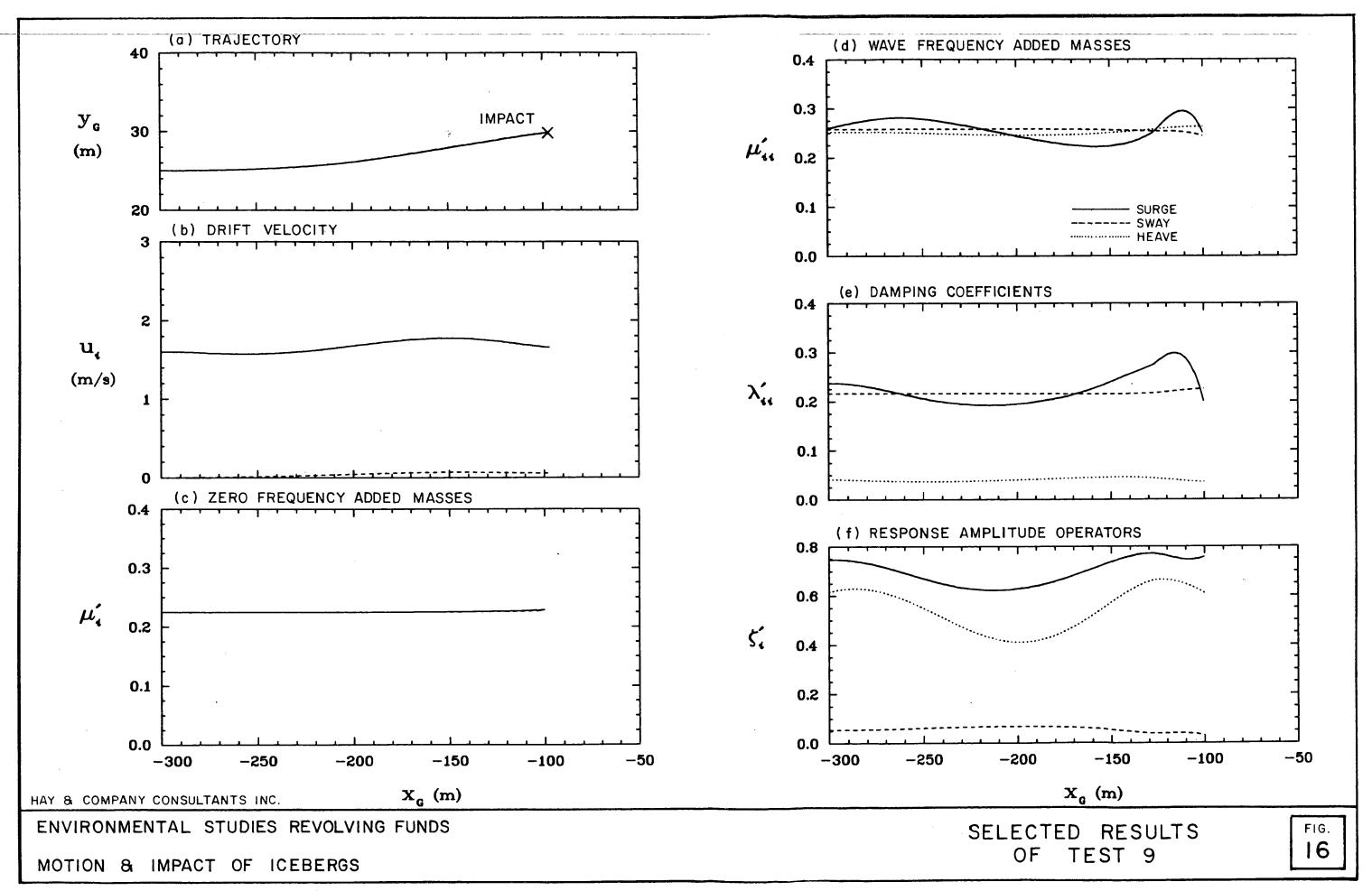
Tests 5-8 described above all refer to the small ice mass B. Tests 9 and 10 refer to the large ice mass A again. Test 9 (Figure 16) corresponds to the same conditions as test 4, with initial offset y = 25 m and wave direction $\alpha = 5$, but now the current velocity is set at 0.4 m/s. In addition, the diffraction and zero frequency added mass calculations are carried out at more frequent intervals, and are performed for the ice mass











closer to impact than in test 4. Consequently, the predicted drift velocity shows a much smoother variation than previously. Furthermore, the hydrodynamic coefficients now show relatively more change just before impact occurs. However, since the facet size is about 25 m, the numerical results obtained immediately prior to impact are unreliable.

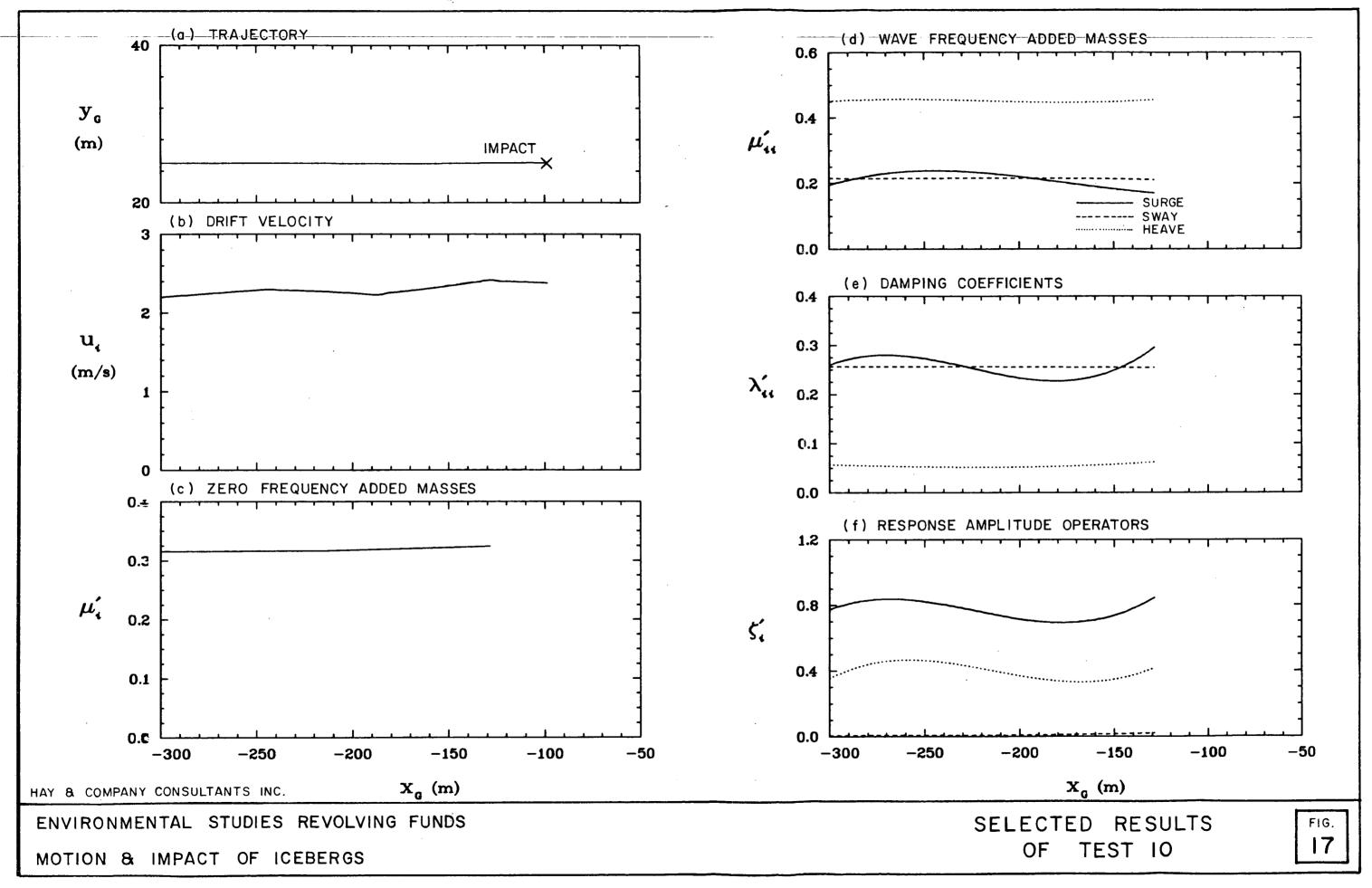
Test 10 (Figure 17) corresponds to the same conditions as test 2, with initial offset y = 25 m and wave direction $\alpha = 0$, except that the water depth is now 60 m, corresponding to a 10 m gap between the bottom of the ice mass and the seabed. One consequence of this change in water depth is that the same wave period gives rise to a shorter wave length (300 m as opposed to 355 m for test 2). Overall, the results of test 10 are not significantly different to those of test 2.

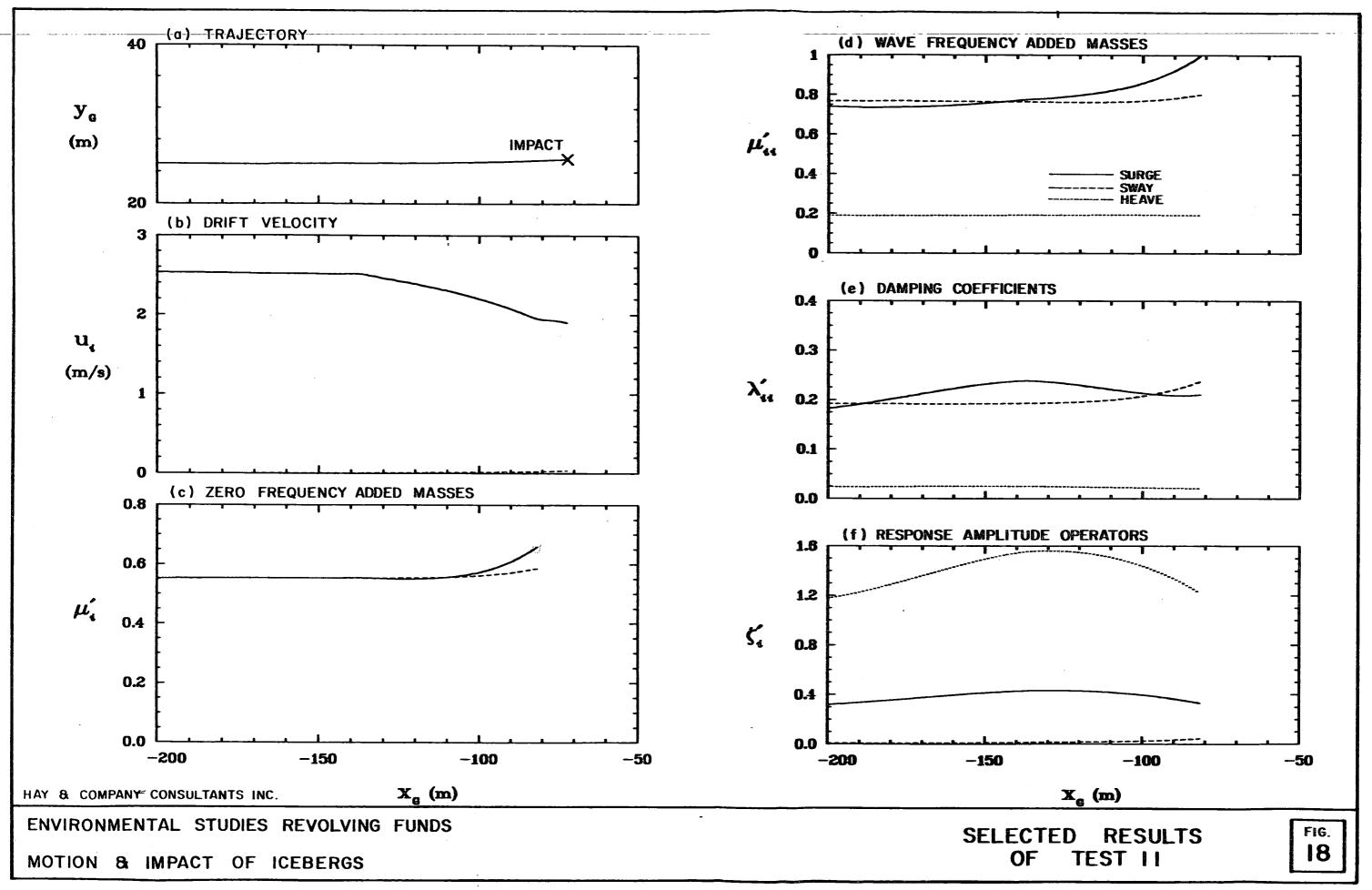
Test 11 (Figure 18) corresponds to the third ice mass considered, ice mass C with a diameter of 50 m, and to an initial offset $y = 25 \, \text{m}$ and wave direction $\alpha = 0$ as in test 2 with ice mass A and test 6 with ice mass B. In this case, the increase in zero frequency added mass prior to impact is quite noticeable. The wave frequency added mass in surge also shows such an increase. The wave drift force remains positive and no drift reversal occurs as in test 6 because ice mass C is significantly larger than ice mass B.

Finally, tests 12 and 13 were carried out with the ice masses A (large) and B (small) respectively, located at specified points far upstream of the structure and with gaps of 5, 10 and 15 m from the structure. The intention here is to use a finer discretization of the ice masses, as indicated in table 3, to investigate the effect of gap width on the zero frequency added masses. Since the degree of discretization alters slightly the numerical values of zero frequency added mass obtained, the results are presented in terms of percentage variations and are given in Table 4 for the large ice mass. The table shows some

increase as the gap is reduced and this eventually falls off for a gap of 5 m as the discretization becomes too coarse for a reliable result. For the small ice mass, the results are not shown since they remain virtually unchanged except that the added mass in sway increases by 1.5% for the 5 m gap. For this case, a relatively small gap is needed to detect an appreciable increase in added mass.

The calculation of the zero frequency added masses involves a Green's function which is calculated by a summation of terms corresponding to multiple reflections of a point source. In order to examine the effect of a small gap more efficiently, a more efficient routine should be developed for the Green's function. This is possible by applying algebraic manipulation to the infinite series expression for this function. A theoretical study of the zero frequency added mass immediately prior to and after impact would be appropriate in this context.





8.0 SURVEY OF MODEL TANK FACILITIES

Part of the terms of reference for this study is a survey of laboratory facilities capable of undertaking a model study of iceberg drift in the vicinity of an offshore structure. The survey was undertaken in two parts:

- Identify facilities capable of undertaking studies and obtain data on those facilities
- 2. Evaluate the facilities and make recommendations as to a suitable test program.

While there are innumerable hydraulic laboratories in North America and Europe, the study of iceberg drift in the vicinity of an offshore structure under realistic sea conditions specialized facilities. The need for these specialized facilities narrows the number of potential candidate laboratories considerably. Queries were made to universities, government laboratories, and other institutes to identify candidate facilities. Questionnaires were then addressed to each of laboratories identified while independent research in the open literature was also carried out.

This section of the report describes the results of the survey.

8.1 Description of Required Facilities

Ultimately, the purpose in laboratory testing iceberg drift in the vicinity of an offshore structure is to aid in determining the forces imparted by an iceberg during a collision with a structure. The magnitude of these forces depends on iceberg dimensions; current speed and direction; wave height, period and direction; and initial iceberg offset from the stagnation streamline of the structure. Simulating and controlling all

these variables in a manner which reflects actual open sea conditions requires very special facilities. In choosing the appropriate facilities, the following features should be considered.

A. Facility Size

Model scale effects can be minimized in a facility with large enough dimensions. Basically, there are two types of facilities which may be considered, wave flumes and wave tanks.

Wave flumes are long and narrow and are designed to represent a two dimensional slice of the ocean. capabilities are generally uni-spectral and uni-directional owing to the flume geometry. Most flumes do not create currents by recirculating water but rather tow the model means of a towing carriage (captive model). The advantages of wave flumes are that a large number of facilities are available and the tests tend to be simpler owing to the limited geometry. The disadvantages are that tests are limited to fit the two dimensional geometry (waves and currents in the same direction, etc.) and the size of the structures is limited if blockage effects are to be avoided.

Wave tanks vary in size but differ from flumes in that their horizontal dimensions are comparable. They are designed to test with the full three-dimensional aspects of the open sea. This includes multi-directional and multi-spectral wave capabilities plus currents created by recirculation or by towing. The advantage of wave tanks is the full range of testing possible under simulated ocean conditions while the disadvantage is the complexity of controlling all the test variables.

B. Wave Generating Capabilities

A single location in the open sea can receive waves generated from storm activity located thousands and from several different kilometres away directions simultaneously. In addition, wave spectra from each direction can be different. To truly simulate open conditions, therefore, the facility must be capable of generating multi-directional, multi-spectral waves.

C. Current Generation

There are two possible ways to simulate currents, one is to recirculate the water in the tank and the other is to tow the model with a towing carriage. A uniform current is difficult to achieve by recirculating owing to the small distance available for flow development at the upstream end of the tank and the effect on the velocity profile of the tank boundaries. Towing a model creates a uniform velocity profile as seen by an observer moving with the speed of the carriage. The effect of currents on waves (wave steepening, etc.) is not simulated when towing is used, however, this effect is small for long period waves.

D. Special Considerations

other special considerations are wave reflections from flume boundaries and wave induced currents which non-uniform velocity profiles. In facilities, testing can be carried out for a considerable before these effects become length of time important. size, virtually all facilities Irrespective of utilize beaches to dissipate incident wave energy and these are very effective except for long wave lengths for which the beaches are too small. In addition, some wave generating systems correct for reflections so that the desired

spectrum is maintained. Wave induced currents are more of a problem in small facilities than in large facilities. While the problem is recognized, no corrective measures are taken in the facilities queried in this study beyond an initial mapping of the current field.

Other features which are also important include instrumentation, data aquisition system and, while not a part of the physical facilities, experience in testing floating structures.

8.2 Results of the Survey

The laboratories in the following list were determined to have facilities suitable for testing iceberg drift:

- (1) Maritime Research Institute
 The Netherlands
- (2) Marintek A/S
 Trondheim, Norway
- (3) Swedish Maritime Research Centre Sweden
- (4) National Maritime Institute England
- (5) David Taylor Naval Ship Research and Development Center Bethesda, Maryland
- (6) Hydraulics Laboratory
 Division of Mechanical Engineering
 National Research Council of Canada
 Ottawa, Ontario

- (7) Hydraulic Research Station Wallingford, England
- (8) Hydronautics Ship Model Basin
 Tracor Hydronautics Inc.
 Laurel, Maryland
- (9) Offshore Technology Corporation Escondido, California
- (10) Faculty of Engineering and Applied Science Memorial University of Newfoundland St. John's, Newfoundland
- (11) Institute for Marine Dynamics

 National Research Council of Canada

 St. John's, Newfoundland
- (12) Versuchsanstalt Fur Wasserbav Und Schiffbau Berlin, West Germany
- (13) Danish Ship Research Laboratory
 Lyngby, Denmark
- (14) British Hovercroft Corp. Ltd.

 Experimental and Electronics Labs

 England
- (15) Department of Shipbuilding and Naval Architecture
 University of Strathclyde
 Glasgow, Scotland
- (16) Ocean Engineering Center
 B.C. Research
 Vancouver, B.C.

The laboratories and the particulars regarding their facilities are listed in Table 5. Also included is a rating relative to basin size, wave generating capabilities, current generating capabilities, and experience.

The basis for the rating is:

A. Basin Size

The rating is based on a 40 m x 40 m x 5 m tank being a good size for full testing of three-dimensional aspects of iceberg drift. Facilities significantly larger than these were considered excellent and those smaller were given lesser ratings. Flumes were rated no better than poor owing to the limited testing possible. Where laboratories possess both flumes and tanks, the tanks were reported.

B. Wave Generation

The rating is based on multi-directional and multi-spectral wave generation as being good and the capability of compensating for reflected wave as being excellent.

C. Current Generation

The ability to create currents over the full width and depth of the facility either by towing or recirculation was considered good. The capability of both towing or recirculation was considered excellent.

D. Experience

Experience rating is somewhat more qualitative in nature. It is based on specific questions in the questionnaire.

LABORATORY	1.	PRINCIPAL DIMENSIONS		NSIONS	2.	WAVE GENERATOR					3.	CURRENT CAPABILITIES						4. TOWING CARRIAGE		1	I - EXCELLENT 2 - GOOD	4 - POOR 5 - NON EXISTENT OR				
	ഥംം	WIDTH	MAXIMU	и реттнѕ	NONE	WAVE MAKER	CONTROL	SEA STATE					NONE		CURRENTS						PERCENT UTILIZATION		3 - AVERAGE		NOT TO BE CONSIDERED	
	m	m	Shollow est Area	Deacest Area m		I Piston 2 Poddie	1 Mechanical 2 Electronic 3 Computer	2 UniSpectral WAY 3 MultiSpectral AT (WAV	APPROXIMATE MAXIMUM WAVE HEIGHT, mm AT PERIODS (SECS) of			STRENGTH FULL BASIN	STRENGTH			GRADIENT			O'LEZATION	6. LABORATORY RATING					
						3 Wedge 4 Other			0.5	1.0	2.0	3.0]	DEPTH mm/s	mm/s	Yes	No	Yes	No	Yes No	1	FACILITY SIZE	WAVE GENERATION	CURRENTS	EXPERIENCE	
1. HYDRAULIC RESEARCH LTD WALLINGFORD	25	25	2	6		2	3	3a, 3b	40	150	500	250		200	200	x		х		x	75	3	1	2	,	
	20	20	1.5	2.5		· ,	2	3a, 3b	40	150	350	200		200	200	x		x		x	50	3	1	2	1	
2. SSPD MARITIME RESEARCH AND CONSULTING	88	39	0	3.5		2	2,3	1, 31, 3b			500	.200		1000	1000		x		x	x	100	1	2	1	,	
3. OFFSHORE TECHNOLOGY CORPORATION	42	32	1.5	4.0		1 1	3	3a, 3b		1	459	,		220			x		x	×	60	2	2	2	2	
4. NATIONAL RESEARCH COUNCIL	50	30	0.5	5.7		1.2	3	3a, 3b	40	160	610	740	x				ļi		Ì	×	75	2	1 1	5	3	
5. DANISH MARITIME INSTITUTE	240	12	5.2	5.5		2	2	2 a	40	480	700	850	x					1		x	95	4	3	2	2	
6. MEMORIAL UNIVERSITY	58	4.5	.8	2.1	ļ	2	3	36	ĺ	150	610	480	x					j		x	50	5	3	2	2	
7. B.C. RESEARCH	30	26	2.5	2.5		2	3	1					x					- 1		x	5	2		5		
8. TRACOR HYDRAULIC INC.	127	7.6	-	4		1	3	3a	İ		610		x					l		x		4	3	2	2	
9. VERSUCHSANSTALT FUER WASSERBAU UND SCHIFFBAU	250	8	1	4.8		Pneumatic		За					x	İ				- 1	- 1	x		4	3	2	2	
10. DAVID TAYLOR NSRDC	110	73	6.1	10.7		Pneumatic	2,3	3ъ	}			610	x							x		1	2	2	1	
11. MARITIME RESEARCH INSTITUTE	60	40	1.2	3.1		2	3	3ъ				300		600					İ	×	85	2	2	2	1	
12. HARINTEK A/S	80	50	0	10		2	2,3	3ъ	-	200	600	800		200		x		x		x		1	2	2	1	
13. BRITISH MARITIME TECHNOLOGY	46	30	1.5	2.3		2	2	3a, 3b			400			YES	7	7				x		2	2	2	1	
14. INSTITUTE FOR MARINE DYNAMICS	75	32	3.5	7.5		2	3	3ъ	35	250	1100	600	x			1	l		- 1	x		1	2	5	3	
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The testing envisaged involves releasing a model iceberg in a given sea state and current field and monitoring its trajectory as it approaches an offshore structure. As discussed earlier, two methods of establishing the current field are recirculating flow and towing. Uniform velocity conditions are difficult to achieve in a tank of reasonable dimensions by recirculation. Towing the structure, creates a uniform current field as viewed from the carriage, however, there is some modelling effect on the waves. For long wave lengths, this modelling effect is minimal.

difficulty Whether recirculation towing, the or is in interpreting test results relative to the very complex perhaps non-ideal, input test conditions. Sorting the individual effects of waves, currents, added masses, and other parameters is a problem and the possibility of tank effects results is always a possibility. The best method of sorting out maze is to utilize a computer model in conjunction with testing. Progressive laboratories have adopted this two-model (physical and computer) approach as a research tool and as means of reducing testing costs.

The computer model developed in conjunction with this study makes available a tool for data evaluation which has been absent in previous test programs. This tool lends another dimension to testing which promises more success in the complex process of sorting out the relationship between observed model behaviour relative to measured inputs and determining any model scale effects.

highest rated laboratory from Table 5 is SSPD Research and Consulting in Sweden. Its major advantage laboratories other is that it can create currents The second highest in ranking are recirculation and by towing. David Taylor Naval Ship Research and Development Center the USA and Marintek A/S in Norway. The major disadvantage the the second ranking is that the David Taylor Laboratory can only tow and Marintek can only recirculate to provide currents. A third ranking group includes British Maritime Technology and Hydraulic Research Ltd. in England, and Maritime Research Institute in Holland. The third ranking of these laboratories is mainly due to the size of their facilities.

Owing to the time required in reaching steady state drift under the influence of waves and currents and the difficulty achieving a uniform current field by recirculation, the possibility of undertaking towing tests should be considered in more detail. It is believed that a smaller test tank could be utilized in a towed experiment although the desired sea state is generally only simulated in a small, central area of a small tank. The alternative of recirculation requires a flow development length of approximately ten times the water depth to achieve a uniform velocity distribution.

The choice of model scale is best left up to the discretion of the individual laboratory to optimize the results relative to their facilities. Any of the first three ranked groups possess the facilities and experience to undertake a study. A range of model scales should be suggested, though, to provide an indication of the level of effort.

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APPENDIX: SAMPLE OUTPUT OF THE NUMERICAL MODEL

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175 176	7 15.000	2.199	0.006	1.000	0.0	-267.011	0.061			
177 178 179	11 T 8 17.500	UX 2.199	UY 0.007	UCX 1.000	UCY 0.0	XG -261.515	YG O.080			
180 181 182 183	11 T 9 20.000	UX 2 . 198	UY 0.008	UCX 1.000	UCY 0.0	XG -256.019	YG 0.101		energy pathology of properties	manakana 25
184 185 186 187	11 T 10 22.500	UX 2.198	YU 800,0	UCX 1.000	UCY 0.0	XG -250.523	YG O. 124			
188 189 190	II T 11 25.000	UX 2.198	UY 0.010	UCX 1.000	UCY .	XG -245.028	YG O.180			
192 193 194 195		1	2	3	4	1	5	6		
196 197 198	EXCITING FORCES	(WRT SWL) 0.657	0.060	0.224		0.011	0.113	0.000		
199 200 201 202	MAGNITUDE: PHASE: EXCITING FORCES	15.876	23.552	56.780	*** * * * * * * * * * * * * * * * * * *		165.081	33.231	 scescrate po sapran-1	
203 204 205	MAGNITUDE: PHASE:	0.657 15.876	0.060 23.552	0.224 56.780		9.003 7.718	0.032 19.271	0.000 33.231		
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219 220 221 222	DAMPING COEFFICE									
223 224 225 226	2 3 4	0.202 0.000 0.008 0.000	0.000 0.217 -0.000 0.039 -0.000	-0.000 0.000 0.037 -0.000	-	0.000 0.041 0.000 0.008 0.000	-0.038 -0.000 -0.002 -0.000	-0.000 0.000	·	
227 228 229 230	5 100 000 000 000 000 000 000 000 000 00	-0.035 0.000	-6.000	0.000	<u> </u>	0.000	-0.000	0.000		
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233 234			0.0	0. 0	o.	053	0.004	0.004	0.0	ir. Ni Mula seji		
235 236 237			1	2	3		4	5	6			
238 239 240	RESPONSE	AMPLITUD	E OPERATOR	S (WRT: SWL)								
241 242	MAGNITUD PHASE:		0.661 177.602	0.059 -177.452	0.534 -146.160		0.051 6.565 1	0.615 172.978	0.000 -154.601		erecens Constitution	
243 244 245	RESPONSE	AMPLITUD	E OPERATOR	S (WRT G)								
246 247	MAGNITUD PHASE:		0.526 178.794	0.048 -175.320	0.534 -146.160		0.051 6.565 1	0.615 172.978	0.000 -154.601			:
248 249 250 251 252	WAVE DRI	FT FORCES O	08379	0.00789								
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258 259 260	II - 13	T 30.000	UX 2 . 204	UY 0.013	UCX 1.000	UCY 0.0	XG -234.008	YG 0.215				uar uz e no maaau para a paadaan
261 262 263 264	I1 i4	7 32.500	UX 2.206	UY 0:015	UCX 1.000	UCY O.O	XG -228,490	YG O.254				
265 266 267 268	II B	35.000	UX 2.209	UY 0.016	UCX 1.000	UCY 0.0	XG -222.964	YG O:296				
269 270 271	II 16	T 37.500	UX 2.212	UY 0.018	UCX 1.000	UCY 0.0	XG -217.431	YG 0.342	oruši ir stoki u je nu jedu ji sti	i de National de la Propieta e de Pro	4.5.1.6.80	ATTRICAL STREET OF LICENSE PRO-
272 273 274 275	ZERO FRE		DED MASSES									
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285 286 287 288		T 40.000	2.215	0.019	UCX 1.000	UCY O O	XG -211.891	YG 0.392				
289 290	I I 18	T 42.500	UX 2.218	UY 0.021	UCX 1.000	UCY 0.0	XG -206.342	YG O.447				į

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MAGNI		0.613	0.061	0.154	0.011	0.108	0.000	
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EXCIT	NG FORCE	S (WRT G)						
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	XCITING FORCES	(WRT SWL)					
	AGNITUDE: HASE:	0.828 149.504	0.058 142.237	0.265 -153.896	0.010 142.366	0.153 -30.938	0.000 173.932
	XCITING FORCES	(WRT G)					
	AGNITUDE: HASE:	0.828 149.504	0.058 142.237	0.265 -153.896	0,003 -38.246	0.030 151.773	0.000 173.932
418 419		1	2	3	4	5	.6
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425	2 3	0.000 -0.031	0.254 0.000	+0.000 0.257	0.054	-0.000 0.006	-0.000 -0.000
426 427	4	0.000	0.052	-0.000	0.020	0.000	-0.000
428	5 ????????? # 1983/99	-0.052 0.000	-0.000 -0.000	-0.002 0.000	0.000 -0.000	0.020 -0.000	-0.000 0.000
429 430		0.000	-0.000	0.000	0.000		
431		TENTS (WOT C	wi \				
432 D 433	AMPING COEFFIC	TENIS (WKI S	WL)				
434	1	0.276	-0.001	-0.009	-0.000	-0.052	0.000 -0.000
435 436	2	-0.001 0.003	0.214	0.000 0.044	0.040 -0.000	0.000 -0.001	0.000
437	4	-0.000	0.038	0.000	0.007	0.000	-0.000
438	5	-0.051	0.000	0.002	0.000	0.010 -0.000	-0.000 0.000
439 440	6	0.000	-0.000	0.000	-0.000	-0.000	0.000
441 442			2				
444 445	ISCOUS DAMPING	COEFFICIENT O.O	0.0	0.053	0.004	0.004	0.0
446 447			2	3	4	5	6
448 449							
	ESPONSE AMPLIT	UDE OPERATOR	S (WRT SWL)				
451	ACMITTIDE .	0.773	0.058	0.650	0.052	0.576	0.000
	AGNITUDE:	+55.205	-57.625		114.326	-61.299	-18.955
. 455	ESPONSE AMPLIT	UDESOPERATOR	S (WRIEG)	sertificial contra discrete sal	Bermana Ludaulat dan Historia	ramma a manasi isini maseleo ushili sua	was was widows a sacropped and agreement and substitution that society is a contract of
	AGNITUDE:	0.647	0.047	0.650	0.052	0.576	0.000
458 P	HASE:	-54.009	-55.665	-6.706	114.326	-61.299	-18.955
459 460	AVE DRIFT FORC						
461 W	AVE DRIFT FORC	ES					
462 463		0.05831	0.00335				
463 464							

Listing of 6			1 JUL 7, 19 DED MASSES	85 for CC1	J≐MD\$Q Page								
466 467	zko (r k c	1	0.229	-0.000	-0.003		-0.000	-0.048	0.000	an a ser in a la serie	 	•••	
468 469 470		2 3 4	-0.000 0.000 -0.000	0.228 0.000 0.045	0.000 0.727 0.000	ta dare tarretti	0.047 0.000 0.016	0.000 -0.000 0.000	-0.000 0.000 -0.000				
471 472		5 6	-0.045 0.000	0.000 -0.000	0.001 0.000		0.000 -0.000	0.016 -0.000	-0.000 0.000				
473 474 475						UVILENI							
476 477	11 32	77.500	UX 2.352	UY 0:050	UCX 1.000	UCY O.O		YG 1.783					
478 479 480	11	т	UX	UY	UCX	UCY	ХG	YG					
481 482	33	80.000	2.343	0.050	1.000	0.0	-120.177	1.907					
483 484 485	1 I 34	T 82.500	UX 2.333	UY 0.049	UCX 1.000	UCY O.O	XG - 1 14 . 357	YG 2.030		ica i se secul sellibracións			2972353
486 487 488	11	T	UX	UY	ucx	UCY	. XG	YG					
489 490	35	85.000	2.322	0.049	1.000	0.0						, and the second	
491 492 493	I I	T 87 500	UX 2.309	UY 0.048	UCX 1.000	UCY 0.0		YG 2.269					
494		MPACT **		N TERMINATI									